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Abstract: This paper demonstrates that geochemistry is useful for the identification of sediment origin and provenance in alluvial plains characterised by a complex hydrographic evolution. The study is focused on the northeastern Padanian plain (Italy), an area primarily characterised by sedimentary contributions from the two largest Italian river systems (Po and Adige), which intimately interacted during the last millennia. X-ray fluorescence analyses of 120 soils and alluvial sediments define three diverse geochemical affinities that have distinctive siderophile/chalcophile elemental ratios. The sample group characterised by high Ni/Zn and Cr/Pb values conforms to modern Po River sediments, whereas a second group showing low Ni/Zn and Cr/Pb values conforms to the geochemical signature of modern Adige River sediments. A third sample group defines a "transitional" affinity that represents a geochemical mixture of Po (70%) and Adige (30%) sedimentary end-members. Based on these geochemical features, it is possible to distinguish alluvial sediments of the Po River basin ($Ni/Zn > 1.0$ and $Cr/Pb > 4.2$) from those of the Adige River basin ($Ni/Zn < 0.6$ and $Cr/Pb < 1.9$) and to provide evidence of the migration of these rivers during the evolution of the Padanian plain. The interpretation of transitional samples is less constrained and could imply an ancient connection between the two fluvial systems, possibly due to the development of wetlands where both the Po and Adige rivers variably delivered their sedimentary contributions. This study approach, therefore, provides important implications for palaeohydrographic and palaeoenvironmental reconstructions in a complex area that is characterised by significant geomorphological modifications during the last millennia.

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Geochemical proxies of sediment provenance in alluvial plains with interfering fluvial systems: ~~a~~A study case from NE Italy

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Abstract

This ~~contribution paper~~ demonstrates that geochemistry is useful for the identification of sediment origin and provenance in alluvial plains characterised by a complex hydrographic evolution. The study is focused on the northeastern Padanian plain (Italy), an area ~~primarily~~mainly characterised by sedimentary contributions from the two largest Italian river systems (Po and Adige), ~~which that~~ intimately interacted ~~during~~in the last millennia. X-ray fluorescence analyses of 120 soils and alluvial sediments define three diverse geochemical affinities ~~having that~~ have distinctive siderophile/chalcophile elemental ratios. The sample group characterised by high Ni/Zn and Cr/Pb values conforms to ~~the~~ modern Po ~~river~~River sediments, whereas a second group showing low Ni/Zn and Cr/Pb values conforms to the geochemical signature of ~~the~~ modern Adige ~~river~~River sediments. A third sample group defines a “transitional” affinity ~~representing that~~ represents a geochemical mixture of Po (70%) and Adige (30%) sedimentary end-members. Based on these geochemical features, it is possible to distinguish alluvial sediments of the Po ~~river~~River basin (Ni/Zn > 1.0 and Cr/Pb > 4.2) from those of the Adige ~~river~~River basin (Ni/Zn < 0.6 and Cr/Pb < 1.9) and to provide ~~evidence of~~elues on the migration of these rivers during the evolution of the Padanian plain. The interpretation of transitional samples is less constrained and could imply an ancient connection between the two fluvial systems, ~~possibly due to~~plausibly by the development of wetlands where both ~~the~~ Po and Adige rivers variably delivered their sedimentary contributions. This study approach, therefore, provides important implications for palaeohydrographic and palaeoenvironmental reconstructions in a complex area that ~~was~~is characterised by significant geomorphological modifications ~~during~~in the last millennia.

Keywords: Padanian plain; Po and Adige rivers; geochemical proxies; palaeohydrographic reconstruction

Introduction

The sedimentary infill of continental basins is related to climatic and tectonic processes that influence rock weathering, erosion, transport of clastic particles and deposition. The interplay of varying subsidence rates and sediment supplies results in the development of different sedimentary environments, ranging from fluvial to a variety of lacustrine settings. The resulting floodplains are complex landscape assemblages ~~of landscapes~~ composed of active landforms (e.g., channels) and stabilised units of various ages and origins (e.g., abandoned channels and banks), which are often ~~formed~~derived by the overlapping ~~in space and time~~ of multiple riverine systems in space and time.

38 The development of human ~~settlements~~ settings and their interactions with the hydrographic
39 networks ~~causes~~ provides additional complexities ~~into~~ the natural evolution of these environments.
40 Floodplains are, therefore, paleoclimatic and ~~hystorical~~ historical archives; ~~that recording human~~
41 ~~activities and the~~ hydroclimatic evolution, ~~and human activities that which~~ are ~~typically~~ usually
42 ~~revealed by the~~ deciphered coupling of geomorphological, sedimentological and archaeological
43 studies (Gautier et al., 2009, Hoffmann et al., 2010; D’Haen et al., 2012). In this framework, the
44 ~~understanding of~~ sedimentary dynamics in alluvial plains can be ~~determined~~ implemented by the
45 ~~usage of~~ geochemical tracers that ~~are~~ represent powerful tools for the identification of ~~the~~ sediment
46 source areas (e.g., Macklin et al., 2006; Singh, 2010; Campodonico et al., 2016). ~~This~~ approach
47 has been widely used in the Emilia-Romagna region ~~in order~~ to distinguish the alluvial sediments
48 ~~in of the~~ Po ~~river~~ River, ~~which draining~~ the Western Alps, from those of other riverine systems ~~that~~
49 ~~draining~~ the Northern Apennines (Amorosi, 2012; Bianchini et al., 2012; 2013). Conversely, the
50 potential of ~~this~~ method has ~~been~~ scarcely ~~been~~ tested northward, where the Po ~~river~~ River system
51 interfered with other fluvial depositional systems during the ~~development of the~~ Padanian plain
52 ~~development~~. ~~In particular, in~~ this work we particularly investigated a sector of the plain, located
53 between the towns of Ferrara and Rovigo, that is bordered to the northward and southward by the
54 Adige and Po rivers, respectively (Fig. 1). The purpose was to identify geochemical proxies for
55 distinct sedimentary end-members and to analyse their spatial distribution, ultimately providing
56 ~~evidence~~ clues for the palaeohydrographic evolution of the area. ~~To achieve this~~ ~~in this view~~, 120
57 major and trace element geochemical analyses of soils and alluvial sediments have been statistically
58 ~~studied~~ elaborated in order to identify the geochemical trends related to weathering of different
59 mother rocks ~~that~~ outcropping in separate hydrological basins. Moreover, 18 samples of bed and
60 suspended sediments from the Po and Adige rivers have been ~~also~~ sampled and analysed, ~~in order~~
61 to verify their geochemical affinity and to compare the current and past ~~daily~~ trends ~~with those~~
62 ~~occurred in the past~~. The results represent a powerful complementary tool to the geomorphological
63 approach for the reconstruction of the palaeohydrographic patterns of the northeastern Padanian
64 plain, ~~which that~~ are currently ~~so far~~ hypothetical and primarily ~~mainly~~ based on the interpretation of
65 landforms and ~~findings of~~ archaeological remains (Piovan et al., 2010; Corrà and Mozzi, 2016).

66

67 **Geological and geomorphological outlines**

68 The study area is located in the easternmost part of the Padanian plain in northern Italy, a
69 sedimentary basin between the Alps and the Apennines that was progressively filled by Neogene
70 marine sediments and finally capped by fluvial sediments since the Pleistocene. ~~During the~~ ~~In~~ Late
71 Holocene, the sedimentary discharge was supplied by two primary ~~main~~ fluvial systems (Po and
72 Adige), and the configuration of the local drainage system was predominantly ~~mainly~~ controlled by
73 the interaction between ~~among~~ tectonic, climatic, eustatic, and sedimentary processes (Garzanti et
74 al., 2011). The geological features of the ~~is~~ basins are briefly described below.

75 The Po ~~river~~ River drains the western Alps (Marchina et al., 2015), ~~which that~~ are characterised by a
76 prevalent crystalline basement, consisting of felsic plutonic rocks (granitoids) and metamorphic
77 rocks (Boriani et al., 2003). Jurassic ophiolites consisting of serpentinized peridotites, gabbros and
78 basalts are also observed in the western Alps, together with sedimentary rocks such as Mesozoic
79 (Triassic and Jurassic) limestones that become significant only in the central Alps (in Lombardy;

80 Bosellini, 2004) and subordinate Tertiary silico-clastic rocks (Maino et al., 2013). Ore deposits
81 recorded in the Po basin ~~primarily~~mainly consist of Fe-Ni and PGE (platinoids) elements,
82 typicallyusually associated ~~with~~ ultramafic rocks (Garuti and Rinaldi, 1986; Rossetti and
83 Zucchetti, 1988; Carbonin et al., 2015).

84 The Adige ~~river~~River drains metamorphic rocks of the Austroalpine domain in the upper part of the
85 basin, rhyolitic rocks of the Atesian volcanic complex between Bolzano and Trento, and ~~then~~ the
86 carbonate units of the Southalpine domain (Natali et al., 2016). Zn-Pb-Cu-Fe sulphide-bearing ore
87 deposits are renowned in the Adige basin (Nimis et al., 2012), ~~which are occasionally~~sometimes
88 associated with barite and fluorite (Bakos et al., 1972).

89 The study area ranges in elevation between 5 and 3.5 m a.s.l. and is characterised by a complex
90 network of alluvial ridges formed by the aggradation of sandy and silty channel deposits, natural
91 levees, and minor, proximal crevasse splays. The depressions contain finer distal levee deposits,
92 such as silty clays, ~~that are sometimes~~occasionally associated with abundant organic matter,
93 suggesting the local development of swamps (Piovan et al., 2008; 2010).

94 The soils developed on these alluvial deposits are generally characterised by a limited profile
95 development, in which the lack of soil maturity is related to the young depositional age (Late
96 Holocene), fluvial reworking and extensive agricultural activities (ploughing). These young soils,
97 therefore, preserve the geochemical signature of ~~the~~ alluvial sediments from which they ~~are~~ derived,
98 ~~which~~ in turn ~~are~~ related to ~~the nature of~~ the parent rocks ~~that~~ outcropping in the various
99 hydrological basins that delivered sediments to the alluvial plain.

100

101 **Materials and methods**

102 Alluvial soils and sediments were collected from 60 sampling sites distributed in 5 distinct
103 agricultural areas of the Rovigo province (Boara Pisani, Concadirame, Lusia, Fratta Polesine-
104 Bosaro, Stienta), located between ~~the~~ Po and Adige rivers (Fig. 1). According to the 1:250,000 soil
105 map of the Veneto region (Osservatorio Regionale Suolo dell'ARPAV, 2015), the investigated soils
106 are ~~primarily~~mainly Cambisols, subordinately Fluvisols and rarely Calcisols. Following the
107 sampling strategy adopted by Bianchini et al. (2012) and (2013), two samples were collected by an
108 Eijkelpamp hand auger at each ~~selected~~ site: one representative ~~sample~~ of the soils just beneath the
109 roots zone (at a depth of 20-30 cm) and another ~~one~~ representative ~~sample~~ of the underlying alluvial
110 sediments (at a depth of 90-110 cm). ~~Taking into e~~Considering that the ~~tillage~~ depth ~~of tillage~~
111 ~~for the~~involving soil digging, stirring, and overturning is ~~typically~~usually ca 50 cm, the shallower
112 samples are considered representative of the soil portion ~~affected~~interested by agricultural activity,
113 while the deepest samples are interpreted as representative of horizons unaffected by anthropogenic
114 activities. Samples were air-dried and divided into distinct aliquots for grain-size and geochemical
115 investigations. The grain size determination was obtained on a sample subset, following the
116 notional classification of Wentworth, by sieving the sandy from the fine (<63 µm) fraction, and then
117 using wet gravitational separation in deionised water to ~~separated~~divide the clay ~~and from the~~ silt
118 fractions ~~using~~by a "Micrometrics Sedigraph 5100".

119 | The amount of ~~S~~soil ~~O~~rganic ~~M~~atter (SOM) was estimated through thermogravimetry performed
120 | at 500°C and expressed as percent weight (wt%) loss with respect to the starting weight of the
121 | material (Nelson and Sommers, 1996; Abella and Zimmer, 2007).

122 | ~~For As concerns~~ the modern river sediments, sampling was ~~performed~~~~carried out~~ following the
123 | ~~method~~~~approach~~ described by Corazzari et al. (2016). The solid suspended load of ~~the~~ Adige and Po
124 | rivers was obtained by collecting 10 litres of river water in plastic containers; after ~~the~~ particles
125 | ~~settling~~, the supernatant liquid was removed ~~using~~~~by~~ a siphon, and the residue was dried in an
126 | oven at 60°C. Bed sediments (ca. 1 kg) were collected from the river bank and dried in an oven at
127 | 60°C.

128 | For the geochemical analysis, the samples were powdered in ~~an~~ agate mill and ~~approximately~~~~about~~
129 | 4 g of powder was ~~hydraulically~~ pressed with ~~addition of~~ boric acid ~~by hydraulic press~~ to obtain
130 | powder pellets. Simultaneously, a sample aliquot of 0.5–0.6 g was heated for ~~approximately~~~~about~~
131 | 12 h in a furnace at ~~temperature of~~ 1000 °C ~~in order~~ to determine the ~~L~~oss ~~O~~n ~~f~~ignition (LOI).
132 | This parameter measures the total concentration of volatile species contained in the sample. The
133 | Wavelength Dispersive X-Ray Spectrometry (WDXRF) analysis of the powder pellets was
134 | ~~performed~~~~carried out~~ using an ARL Advant'X spectrometer Thermo Scientific (Waltham, MA,
135 | USA). Calibrations were obtained analysing certified reference materials, and ~~a~~ matrix correction
136 | was performed according to the method proposed by Trail and Lachance (1966). Precision and
137 | accuracy calculated by repeated analyses of international standards (with matrices comparable
138 | ~~to~~~~with~~ the studied samples, Di Giuseppe et al., 2014) were generally better than 3% for Si, Ti, Fe,
139 | Ca and K, and 7% for Mg, Al, Mn and Na. For ~~the~~ trace elements (above 10 ppm), the errors were
140 | generally better than 10%.

141 | Geostatistical modelling was ~~performed~~~~carried out in order~~ to show the spatial variation of
142 | geochemical data. Interpolated maps along the geographical extension were produced in ArcGIS 9.3
143 | (Geostatistical Analyst extension) at a resolution of 100 m ~~using~~ ~~aby~~ generalised linear regression
144 | technique (ordinary kriging), ~~using~~ ~~and~~ a spherical semivariogram model with ~~the~~ nugget on log₁₀-
145 | transformed data.

146 |

147 | **Results**

148 | The bulk rock geochemical composition of 120 soil and sediment samples from the alluvial sector
149 | located between the Adige and Po rivers are presented in Supplementary Table 1-5, and
150 | summarised in Table 1. For most samples, the textural composition, ~~is reported~~ based on the
151 | classification scheme of Wentworth, ~~is also reported~~.

152 | The WDXRF and textural analyses reveal that the collected samples span ~~over~~ a wide range of
153 | chemical and grain size compositions. ~~In particular~~ ~~for~~ ~~the~~ major elements, SiO₂ varies from 35.70
154 | ~~up~~ to 72.79 wt%, Al₂O₃ from 8.40 ~~up~~ to 17.24 wt%, CaO from 1.09 ~~up~~ to 11.56 wt% and Fe₂O₃
155 | from 2.95 ~~up~~ to 10.56 wt%. The grain size distribution also shows ~~significant~~~~a great~~ variability
156 | ~~for~~~~with~~ ~~the~~ sand and clay contents varying in the ranges ~~of~~ 0.1 - 92.4 % and 3.0 - 65.4 %,
157 | respectively, whereas ~~the~~ SOM varies between 0.7 and 24.8. The sand content (%) is positively

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Comment [QCE2]: Please be consistent with the inclusion of a space between the number and percent symbol.

158 | correlated with SiO₂ (wt%), whereas the clay and SOM contents (%) are positively correlated with
159 | Fe₂O₃, K₂O and Al₂O₃ (wt%).

160 | For each sampling site, ~~the~~ comparison between the metal concentration at 20-30 and 90-110 cm in
161 | depth ~~was performed to~~ allowed the evaluation ~~the of~~ possible T_{top} E_{enrichment} F_{factors} (TEF)
162 | induced by anthropogenic activities. ~~In this view, t~~The trace element budget of the studied
163 | soils/sediments preserves a “geogenic” signature as demonstrated by the general lack of TEFs
164 | (Trevisan et al., 2016). The heavy metal distribution highlights different backgrounds in the distinct
165 | sampling areas: Stienta and ~~to a~~ lesser extent Bosaro-Fratta Polesine ~~show comparative~~
166 | relatively enriched ~~of in~~ siderophile elements (Cr, Ni, Co, V) and depleted ~~in~~ in chalcophile
167 | elements (Zn, Cu, Pb), whereas the opposite is observed at Boara Pisani, Concadirame and Lusia.
168 | These observations, ~~shown~~ highlighted by box-plots (Trevisan et al., 2016), are emphasised in the
169 | geochemical maps (Fig. 2), which show a northward increase of Zn and Pb (chalcophile elements)
170 | and a southward enrichment in Ni and Cr (siderophile elements) in the alluvial sediments of this
171 | sector of the Padanian plain, suggesting that different geochemical affinities characterise the Po and
172 | Adige river basins. Discrimination diagrams, useful to identify the provenance of the investigated
173 | samples, are reported in Fig. 3. The Ni vs Zn diagram (Fig. 3a) ~~allows~~ identifies ~~two~~ data series
174 | ~~that possibly~~ plausibly ~~indicate~~ referred to the sedimentary contributions from the Po and Adige
175 | rivers, respectively, whereas a third sample group delineates a “transitional” series with an
176 | intermediate character. ~~In particular, s~~Sediments from the Po ~~river~~ River basin (45 samples, r² =
177 | 0.67) are characterised by the highest Ni/Zn values (average 1.78, SD = 0.28), whereas sediments
178 | from the Adige ~~river~~ River basin (49 samples, r² = 0.56) ~~have~~ by the lowest Ni/Zn values (average
179 | 0.38, SD = 0.09); the transitional series also displays a very good correlation between Ni and Zn (28
180 | samples, r² = 0.79) and intermediate Ni/Zn values (average 0.74, SD = 0.13). Coherently, the Ni/Zn
181 | values of the sediments attributed to the Po river closely correspond to those characteristic of the Po
182 | sediments (average 1.81, SD = 0.37) ~~that~~ outcropping southward in the Province of Ferrara
183 | (Bianchini et al., 2013). The discriminated series are also characterised by other distinctive
184 | siderophile/chalcophile values such as Cr/Pb (Fig. 3b), which is highest in the Po (average 7.95, SD
185 | = 2.1), lowest in the Adige (average 1.41, SD = 0.35) and intermediate in the transitional (average
186 | 2.39, SD = 0.75) sediments. ~~The existence of t~~Three sample groups ~~are~~ is also ~~shown~~ highlighted
187 | ~~in~~ by the Ba vs Al₂O₃ diagram (Fig. 3c); that indicates a general enrichment of Ba in the fine
188 | fraction, and comparatively higher values in the Adige alluvial sediments with respect to those from
189 | the Po sediments. The different concentrations of siderophile/chalcophile elements in the recognised
190 | sample groups is generalised in the Cr+Ni+Co+V vs Zn+Pb+Cu+Ba diagram (Fig. 4a) and
191 | schematised in the Ni/Zn vs Cr/Pb diagram (Fig. 4b) that ~~provide~~ allow a geochemical
192 | classification scheme for the sediment ~~provenance of sediments~~ in the investigated area.
193 | ~~Discrimination is also provided by t~~The relationships between alkaline and alkaline-earth elements
194 | ~~are also used to distinguish the sediments~~: Po sediments are ~~relatively~~ comparatively enriched in Sr
195 | (ppm), whereas the Adige sediments are ~~relatively~~ comparatively enriched in K₂O (wt%) and Ba
196 | (ppm). Further geochemical differences are ~~shown~~ given by P₂O₅ and MnO (wt%), ~~which~~ that are
197 | generally higher in the Adige and Po samples, respectively.

198 |
199 | The bulk rock geochemical composition of 18 bed and suspended sediments from the Po and Adige
200 | rivers are presented in Supplementary Table 6, and summarised in Table 2. The major element
201 | composition of the bed sediments is, in general, ~~relatively~~ comparatively enriched in SiO₂ compared

202 ~~to the~~with respect to suspended sediments, ~~which~~that are ~~in turn~~ enriched in Al₂O₃ and K₂O (wt%).
203 This observation reflects the prevalence ~~of~~ presence of quartz in the coarser bed sediments and of clay
204 minerals and carbonates in the finer suspended particles. ~~The B~~bed sediments are ~~in~~ generally
205 depleted in Ni, V, Zn, Pb, and Cu ~~compared with respect~~ to ~~the~~ suspended sediments, ~~which~~
206 ~~indicates~~ing that most heavy metals are preferentially partitioned in clay minerals and/or in the
207 associated amorphous matter. ~~A~~The comparison between sediments from ~~the~~ Po and Adige rivers
208 ~~shows~~reveals that the former ~~is~~are ~~relatively~~comparatively enriched in CaO ~~compared with respect~~
209 to MgO (wt%). ~~For the~~As concerns heavy metals, the Po ~~river-River~~ sediments are characterised by
210 higher concentrations of siderophile elements whereas the Adige ~~river-River~~ sediments are
211 comparatively enriched in chalcophile elements, ~~showing with~~ geochemical trends overlapping
212 those of their alluvial counterparts (Figs. 3 and 4). Coherently, ~~the~~ Po ~~river-River~~ sediments
213 ~~have~~show higher Ni/Zn (average 1.46, SD = 0.28) and Cr/Pb (average 5.36, SD = 1.10) ~~ratios~~
214 ~~compared with respect~~ to ~~the~~ Adige ~~river-River~~ sediments (average Ni/Zn = 0.32, SD = 0.04;
215 average Cr/Pb = 1.21, SD = 0.38). ~~The~~ Po sediments ~~have~~show a relatively high ~~Sr~~ content ~~of Sr~~
216 (ppm), whereas ~~the~~ Adige sediments are characterised by a comparative ~~Ba~~ enrichment ~~of Ba~~
217 (ppm).

218 Discussion

220 The reported geochemical data ~~for~~on alluvial sediments from the northeastern Padanian plain
221 provide constraints for the Holocene palaeohydrographic evolution of the area, ~~which~~that is
222 ~~primarily~~mainly determined ~~known~~ by geomorphological studies (often based on remote sensing
223 data), archaeological ~~discoveries~~findings and historical chronicles (e.g., Castiglioni, 1999).

224 The major element composition primarily reflects the mineralogical features of distinct depositional
225 facies. The positive correlation between sand content (%) and SiO₂ (wt%) suggests that the coarse
226 fraction (channel deposits) is generally dominated by quartz and feldspars, whereas the increase of
227 clay and SOM contents in parallel with Fe₂O₃, K₂O and Al₂O₃ (wt%) suggests that the fine fraction
228 (distal levee deposits) includes iron-rich clay minerals, such as chlorite and smectite, and that it is
229 intimately associated with the organic matter. However, major elements are scarcely discriminant
230 for the sediment provenance in the study area, where the Po and Adige river systems ~~variably~~ously
231 migrated and interfingered during the last millennia. ~~In fact, a~~Among ~~the~~ major elements, only CaO
232 and MgO (wt%) give ~~an approximate~~rough indication ~~for~~about the sediment provenance, as the
233 CaO/MgO ratio is generally higher in the Po ~~river-River~~ deposits.

234 ~~In contrast, On the contrary, in the~~ for the current case study, we have demonstrated that the trace
235 element signature is a precise tool for the identification of the Po and Adige end-members, ~~which~~
236 ~~refines~~ing the discrimination criteria previously proposed by Amorosi et al. (2008) and Picone et al.
237 (2008). According to our inferences, alluvial sediments with Ni/Zn > 1.0 and Cr/Pb > 4.2 are clearly
238 attributable to the Po ~~river-River~~ basin. This geochemical fingerprint, characterised by a relative
239 enrichment in siderophile elements, is related to the presence of mafic and ultramafic rocks
240 (~~primarily~~mainly included in ophiolite sequences) ~~that~~ outcropping within the Po ~~river-River~~
241 hydrological basin (Amorosi, 2012; Bianchini et al., 2012; 2013). ~~Conversely, On the other hand,~~
242 alluvial sediments with Ni/Zn < 0.6 and Cr/Pb < 1.9 are clearly attributable to the Adige ~~river-River~~
243 basin, where the relative enrichment of chalcophile elements is related to the presence of Zn-Pb-Cu

244 sulphide ore deposits, ~~which are occasionally sometimes~~ associated with barite in the mountainous
245 part of the catchment. These discrimination criteria have been validated ~~by with the Po and Adige~~
246 ~~river sediments of Po and Adige data~~ retrieved from the available literature (Boldrin et al., 1989;
247 Dinelli and Lucchini, 1999). ~~The~~ Po soils and sediments are ~~relatively comparatively~~ enriched in Sr
248 (ppm) resulting from the weathering of plagioclase-bearing mafic rocks, in addition to the
249 geochemical budget related to carbonates. Conversely, ~~the~~ Ba (ppm) enrichment of ~~the~~ Adige
250 sediments is possibly related to the weathering of alkaline feldspar-bearing felsic rocks and/or to
251 barite deposits. Moreover, ~~the~~ Adige sediments are generally enriched in P₂O₅ (wt%) whereas ~~the~~
252 Po sediments ~~are enriched~~ in MnO (wt%), ~~whereas. However,~~ the distribution of other elements,
253 such as Zr and REEs, ~~doesn't~~ not differentiate the sediments of the two riverine systems.

254 Samples having transitional characteristics between the Po and Adige geochemical affinities
255 represent a mixture of the above described end-members. They could result from the juxtaposition
256 of several alluvial events ~~foref~~ the two fluvial vectors occurring ~~in~~ in historical times when both ~~the~~
257 Po and Adige ~~rivers~~ were scarcely channelled, thus more frequently subjected to overflow and
258 avulsion (Stefani and Vincenzi, 2005). However, remarkable inter-element correlations (Figs. 3 and
259 4) suggest that the transitional series represents a ~~quite~~ precise mixture of ~~the~~ Po and Adige
260 sediments in the proportions of 70% and 30%, respectively. This homogeneity implies the existence
261 of an ancient fluvial system collecting water and sediments from both river basins, ~~and~~ flowing
262 between the current stream paths of ~~the~~ Po and Adige rivers, ~~which is possibly plausibly~~ represented
263 by a series of palaeochannels known as “Po di Fratta”, “Cona-Saline” (Bronze age) and “Po di
264 Adria” (Iron age-Roman times; Piovan et al., 2010; 2012; Ravazzi et al., 2013). The existence of
265 these ancient branches of the Po ~~river-River~~ has been confirmed by petrographic studies on coarse-
266 grained sediments from alluvial ridges near Rovigo, that highlighted the presence of key minerals
267 (glaucofane and serpentine) attributable to the weathering of rocks (blueschists and serpentinites)
268 ~~that~~ outcropping exclusively in the Po ~~river-River~~ basin. Unfortunately, the approach adopted by
269 previous studies is appropriate only for palaeochannel deposits, whereas ~~it~~ cannot be applied to
270 finer sediments, such as those of distal levee facies, ~~which-that~~ can ~~only~~ be interpreted ~~using only by~~
271 bulk sample geochemistry. Our new data suggests that this ancient Po ~~river-River~~ stream path
272 possibly captured a branch of the Adige fluvial system (also including minor distributaries oriented
273 southward, Castiglioni, 1999), thus acquiring part of its sedimentary load that effectively mixed and
274 generated the transitional sediments widespread in the study area. However, this hypothesis is
275 scarcely constrained by the available geomorphological evidence (Piovan et al., 2012; Fontana et
276 al., 2014) and historical chronicles (Calzolari, 2008); and, therefore, it is more plausible that the
277 interaction between the Po and Adige fluvial systems (that ~~variably~~ ~~interfingered~~ ~~interfered~~ in
278 space and time ~~during~~ the last millennia) occurred through the development of wetlands fed by
279 both ~~the~~ Po and Adige rivers, especially in Pre-Roman times (Stefani and Vincenzi, 2005). This
280 interpretation is supported by the prevalent fine texture of samples belonging to the transitional
281 series, as well as by historical chronicles ~~that which~~ describe the widespread presence of flooded
282 zones in the studied area (Whitehouse, 2001).

283

284 Conclusions

285 The geochemical fingerprint of the ~~sediments from the~~ Province of Rovigo sediments clearly
286 characterises/reflects the affinity of the solid load transported by the two primary/main riverine
287 systems that envelope the area, i.e., the Po ~~river-River containing/having~~ sediments characterised
288 by notable concentrations of Ni and Cr, and the Adige ~~river-River containing/having~~ sediments
289 relatively/comparatively enriched in chalcophile elements such as Pb and Zn. Therefore,
290 Ggeochemistry ~~therefore~~ represents a fundamental tool to correlate/relate the investigated alluvial
291 soils and sediments to a specific basin characterised by distinct parent rocks. ~~In this view,~~
292 gGeochemical maps showing/reporting the spatial distribution of key elements are useful to define
293 geochemical gradients and local anomalies, but the precise identification of the sediment
294 provenance can only be/only assessed ~~by the~~ usage of key elemental ratios. The current study
295 shows/highlights the potential of the geochemical approach in the reconstruction of
296 palaeohydrographic patterns. We demonstrated that the sedimentary influence of ~~Po ~~river-River~~~~
297 sedimentary influence extends northward compared to/respect to its current stream path (as also
298 proposed by authors that studied sediments in the Adriatic Sea; Amorosi et al., 2008; Picone et al.
299 2008), ~~and that/and that~~ the Po ~~river-River~~ sediments are locally present very close to the modern
300 Adige stream path. Moreover, the presented geochemical data suggest that in historical times a
301 branch of the Po ~~river-River~~ crosscutting the studied area was characterised by a sedimentary load
302 with/having a hybrid signature intermediate between those currently recognised in the Po and Adige
303 rivers, ~~nowadays. The existence of t~~These transitional sediments may suggest that at that time the
304 two fluvial systems were connected, i.e., that a branch of the Adige was a tributary of the Po ~~river~~
305 River upstream with respect to the study area or, more probably, that the interference of the two
306 fluvial systems developed through marsh/swamp environments that are recorded in historical
307 chronicles and in geomorphological reconstructions (Vincenzi and Stefani, 2005). These
308 sedimentary facies could, therefore, represent the expression of depositional domains in which both
309 the Po and Adige rivers delivered sedimentary inputs, ~~and were that~~ probably originated within
310 wetlands developed among major fluvial ridges.

311

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313 The authors thank N.P. Branch, an anonymous reviewer and the editor M. Egli for their constructive
314 comments that greatly improved an early version of the paper.

315

316 **Table captions:**

317

318 **Table 1** – Geochemical composition of soils and sediments from the investigated sector of the
319 Padanian plain (Province of Rovigo, NE Italy).

320

321 **Table 2** – Average geochemical composition of modern bed and suspended sediments from the Po
322 and Adige rivers.

323

324

1 **Geochemical proxies of sediment provenance in alluvial plains** 2 **with interfering fluvial systems: A study case from NE Italy**

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5 6 **Abstract**

7 This paper demonstrates that geochemistry is useful for the identification of sediment origin and
8 provenance in alluvial plains characterised by a complex hydrographic evolution. The study is
9 focused on the northeastern Padanian plain (Italy), an area primarily characterised by sedimentary
10 contributions from the two largest Italian river systems (Po and Adige), which intimately interacted
11 during the last millennia. X-ray fluorescence analyses of 120 soils and alluvial sediments define
12 three diverse geochemical affinities that have distinctive siderophile/chalcophile elemental ratios.
13 The sample group characterised by high Ni/Zn and Cr/Pb values conforms to modern Po River
14 sediments, whereas a second group showing low Ni/Zn and Cr/Pb values conforms to the
15 geochemical signature of modern Adige River sediments. A third sample group defines a
16 “transitional” affinity that represents a geochemical mixture of Po (70%) and Adige (30%)
17 sedimentary end-members. Based on these geochemical features, it is possible to distinguish
18 alluvial sediments of the Po River basin ($\text{Ni/Zn} > 1.0$ and $\text{Cr/Pb} > 4.2$) from those of the Adige
19 River basin ($\text{Ni/Zn} < 0.6$ and $\text{Cr/Pb} < 1.9$) and to provide evidence of the migration of these rivers
20 during the evolution of the Padanian plain. The interpretation of transitional samples is less
21 constrained and could imply an ancient connection between the two fluvial systems, possibly due to
22 the development of wetlands where both the Po and Adige rivers variably delivered their
23 sedimentary contributions. This study approach, therefore, provides important implications for
24 palaeohydrographic and palaeoenvironmental reconstructions in a complex area that is characterised
25 by significant geomorphological modifications during the last millennia.

26 **Keywords:** Padanian plain; Po and Adige rivers; geochemical proxies; palaeohydrographic
27 reconstruction

28 29 **Introduction**

30 The sedimentary infill of continental basins is related to climatic and tectonic processes that
31 influence rock weathering, erosion, transport of clastic particles and deposition. The interplay of
32 varying subsidence rates and sediment supplies results in the development of different sedimentary
33 environments, ranging from fluvial to a variety of lacustrine settings. The resulting floodplains are
34 complex landscape assemblages composed of active landforms (e.g., channels) and stabilised units
35 of various ages and origins (e.g., abandoned channels and banks), which are often formed by the
36 overlap of multiple riverine systems in space and time. The development of human settlements and
37 their interactions with the hydrographic networks causes additional complexities in the natural

38 evolution of these environments. Floodplains are, therefore, paleoclimatic and historical archives
39 that record human activities and the hydroclimatic evolution, which are typically revealed by the
40 coupling of geomorphological, sedimentological and archaeological studies (Gautier et al., 2009,
41 Hoffmann et al., 2010; D'Haen et al., 2012). In this framework, the sedimentary dynamics in
42 alluvial plains can be determined using geochemical tracers that are powerful tools for the
43 identification of sediment source areas (e.g., Macklin et al., 2006; Singh, 2010; Campodonico et al.,
44 2016). This approach has been widely used in the Emilia-Romagna region to distinguish the alluvial
45 sediments in the Po River, which drains the Western Alps, from those of other riverine systems that
46 drain the Northern Apennines (Amorosi, 2012; Bianchini et al., 2012; 2013). Conversely, the
47 potential of this method has been scarcely tested northward, where the Po River system interfered
48 with other fluvial depositional systems during the Padanian plain development. In this work we
49 particularly investigated a sector of the plain, located between the towns of Ferrara and Rovigo, that
50 is bordered to the north and south by the Adige and Po rivers, respectively (Fig. 1). The purpose
51 was to identify geochemical proxies for distinct sedimentary end-members and to analyse their
52 spatial distribution, ultimately providing evidence for the palaeohydrographic evolution of the area.
53 To achieve this, 120 major and trace element geochemical analyses of soils and alluvial sediments
54 have been statistically studied to identify the geochemical trends related to weathering of different
55 mother rocks that outcrop in separate hydrological basins. Moreover, 18 samples of bed and
56 suspended sediments from the Po and Adige rivers have been sampled and analysed to verify their
57 geochemical affinity and to compare the current and past daily trends. The results represent a
58 powerful complementary tool to the geomorphological approach for the reconstruction of the
59 palaeohydrographic patterns of the northeastern Padanian plain, which are currently hypothetical
60 and primarily based on the interpretation of landforms and archaeological remains (Piovan et al.,
61 2010; Corrà and Mozzi, 2016).

62

63 **Geological and geomorphological outlines**

64 The study area is located in the easternmost part of the Padanian plain in northern Italy, a
65 sedimentary basin between the Alps and the Apennines that was progressively filled by Neogene
66 marine sediments and finally capped by fluvial sediments since the Pleistocene. During the Late
67 Holocene, the sedimentary discharge was supplied by two primary fluvial systems (Po and Adige),
68 and the configuration of the local drainage system was predominantly controlled by the interaction
69 between tectonic, climatic, eustatic, and sedimentary processes (Garzanti et al., 2011). The
70 geological features of the basins are briefly described below.

71 The Po River drains the western Alps (Marchina et al., 2015), which are characterised by a
72 prevalent crystalline basement consisting of felsic plutonic rocks (granitoids) and metamorphic
73 rocks (Boriani et al., 2003). Jurassic ophiolites consisting of serpentinitised peridotites, gabbros and
74 basalts are also observed in the western Alps, together with sedimentary rocks such as Mesozoic
75 (Triassic and Jurassic) limestones that become significant only in the central Alps (in Lombardy;
76 Bosellini, 2004) and subordinate Tertiary silico-clastic rocks (Maino et al., 2013). Ore deposits
77 recorded in the Po basin primarily consist of Fe-Ni and PGE (platinoids) elements, typically
78 associated with ultramafic rocks (Garuti and Rinaldi, 1986; Rossetti and Zucchetti, 1988; Carbonin
79 et al., 2015).

80 The Adige River drains metamorphic rocks of the Austroalpine domain in the upper part of the
81 basin, rhyolitic rocks of the Atesian volcanic complex between Bolzano and Trento, and the
82 carbonate units of the Southalpine domain (Natali et al., 2016). Zn-Pb-Cu-Fe sulphide-bearing ore
83 deposits are renowned in the Adige basin (Nimis et al., 2012), which are occasionally associated
84 with barite and fluorite (Bakos et al., 1972).

85 The study area ranges in elevation between 5 and 3.5 m a.s.l. and is characterised by a complex
86 network of alluvial ridges formed by the aggradation of sandy and silty channel deposits, natural
87 levees, and minor proximal crevasse splays. The depressions contain finer distal levee deposits,
88 such as silty clays, that are occasionally associated with abundant organic matter suggesting the
89 local development of swamps (Piovan et al., 2008; 2010).

90 The soils developed on these alluvial deposits are generally characterised by a limited profile
91 development, in which the lack of soil maturity is related to the young depositional age (Late
92 Holocene), fluvial reworking and extensive agricultural activities (ploughing). These young soils,
93 therefore, preserve the geochemical signature of the alluvial sediments from which they are derived,
94 which in turn are related to the parent rocks that outcrop in the various hydrological basins that
95 delivered sediments to the alluvial plain.

96

97 **Materials and methods**

98 Alluvial soils and sediments were collected from 60 sampling sites distributed in 5 distinct
99 agricultural areas of the Rovigo province (Boara Pisani, Concadirame, Lusia, Fratta Polesine-
100 Bosaro, Stienta), located between the Po and Adige rivers (Fig. 1). According to the 1:250,000 soil
101 map of the Veneto region (Osservatorio Regionale Suolo dell'ARPAV, 2015), the investigated soils
102 are primarily Cambisols, subordinately Fluvisols and rarely Calcisols. Following the sampling
103 strategy adopted by Bianchini et al. (2012) and (2013), two samples were collected by an
104 Eijkelkamp hand auger at each site: one representative sample of the soils just beneath the root zone
105 (at a depth of 20-30 cm) and another representative sample of the underlying alluvial sediments (at
106 a depth of 90-110 cm). Considering that the tillage depth for the soil digging, stirring, and
107 overturning is typically ca 50 cm, the shallower samples are considered representative of the soil
108 portion affected by agricultural activity, while the deepest samples are interpreted as representative
109 of horizons unaffected by anthropogenic activities. Samples were air-dried and divided into distinct
110 aliquots for grain-size and geochemical investigations. The grain size determination was obtained
111 on a sample subset, following the notional classification of Wentworth, by sieving the sand from the
112 fine (<63 μm) fraction, and then using wet gravitational separation in deionised water to separate
113 the clay and silt fractions using a "Micrometrics Sedigraph 5100".

114 The amount of soil organic matter (SOM) was estimated through thermogravimetry performed at
115 500°C and expressed as percent weight (wt%) loss with respect to the starting weight of the
116 material (Nelson and Sommers, 1996; Abella and Zimmer, 2007).

117 For the modern river sediments, sampling was performed following the method described by
118 Corazzari et al. (2016). The solid suspended load of the Adige and Po rivers was obtained by
119 collecting 10 litres of river water in plastic containers; after the particles settled, the supernatant

120 liquid was removed using a siphon, and the residue was dried in an oven at 60°C. Bed sediments
121 (ca. 1 kg) were collected from the river bank and dried in an oven at 60°C.

122 For the geochemical analysis, the samples were powdered in an agate mill and approximately 4 g of
123 powder was hydraulically pressed with boric acid to obtain powder pellets. Simultaneously, a
124 sample aliquot of 0.5–0.6 g was heated for approximately 12 h in a furnace at 1000 °C to determine
125 the loss on ignition (LOI). This parameter measures the total concentration of volatile species
126 contained in the sample. The Wavelength Dispersive X-Ray Spectrometry (WDXRF) analysis of
127 the powder pellets was performed using an ARL Advant’X spectrometer Thermo Scientific
128 (Waltham, MA, USA). Calibrations were obtained analysing certified reference materials, and a
129 matrix correction was performed according to the method proposed by Trail and Lachance (1966).
130 Precision and accuracy calculated by repeated analyses of international standards (with matrices
131 comparable to the studied samples, Di Giuseppe et al., 2014) were generally better than 3% for Si,
132 Ti, Fe, Ca and K, and 7% for Mg, Al, Mn and Na. For the trace elements (above 10 ppm), the errors
133 were generally better than 10%.

134 Geostatistical modelling was performed to show the spatial variation of geochemical data.
135 Interpolated maps along the geographical extension were produced in ArcGIS 9.3 (Geostatistical
136 Analyst extension) at a resolution of 100 m using a generalised linear regression technique
137 (ordinary kriging) and a spherical semivariogram model with the nugget on log₁₀-transformed data.

138

139 **Results**

140 The bulk rock geochemical composition of 120 soil and sediment samples from the alluvial sector
141 located between the Adige and Po rivers are presented in Supplementary Table 1-5 and summarised
142 in Table 1. For most samples, the textural composition is reported based on the classification
143 scheme of Wentworth.

144 The WDXRF and textural analyses reveal that the collected samples span a wide range of chemical
145 and grain size compositions. For the major elements, SiO₂ varies from 35.70 to 72.79 wt%, Al₂O₃
146 from 8.40 to 17.24 wt%, CaO from 1.09 to 11.56 wt% and Fe₂O₃ from 2.95 to 10.56 wt%. The
147 grain size distribution also shows significant variability for the sand and clay contents varying in the
148 ranges of 0.1 - 92.4 % and 3.0 – 65.4 %, respectively, whereas the SOM varies between 0.7 and
149 24.8. The sand content (%) is positively correlated with SiO₂ (wt%), whereas the clay and SOM
150 contents (%) are positively correlated with Fe₂O₃, K₂O and Al₂O₃ (wt%).

151 For each sampling site, a comparison between the metal concentration at 20-30 and 90-110 cm in
152 depth was performed to evaluate the possible top enrichment factors (TEF) induced by
153 anthropogenic activities. The trace element budget of the studied soils/sediments preserves a
154 “geogenic” signature as demonstrated by the general lack of TEFs (Trevisan et al., 2016). The
155 heavy metal distribution highlights different backgrounds in the distinct sampling areas: Stienta and
156 to a lesser extent Bosaro-Fratta Polesine are relatively enriched in siderophile elements (Cr, Ni, Co,
157 V) and depleted in chalcophile elements (Zn, Cu, Pb), whereas the opposite is observed at Boara
158 Pisani, Concadirame and Lusia. These observations, shown by box-plots (Trevisan et al., 2016), are
159 emphasised in the geochemical maps (Fig. 2), which show a northward increase of Zn and Pb
160 (chalcophile elements) and a southward enrichment in Ni and Cr (siderophile elements) in the

161 alluvial sediments of this sector of the Padanian plain, suggesting that different geochemical
162 affinities characterise the Po and Adige river basins. Discrimination diagrams, useful to identify the
163 provenance of the investigated samples, are reported in Fig. 3. The Ni vs Zn diagram (Fig. 3a)
164 identifies two data series that possibly indicate the sedimentary contributions from the Po and Adige
165 rivers, respectively, whereas a third sample group delineates a “transitional” series with an
166 intermediate character. Sediments from the Po River basin (45 samples, $r^2 = 0.67$) are characterised
167 by the highest Ni/Zn values (average 1.78, SD = 0.28), whereas sediments from the Adige River
168 basin (49 samples, $r^2 = 0.56$) have the lowest Ni/Zn values (average 0.38, SD = 0.09); the
169 transitional series also displays a very good correlation between Ni and Zn (28 samples, $r^2 = 0.79$)
170 and intermediate Ni/Zn values (average 0.74, SD = 0.13). Coherently, the Ni/Zn values of the
171 sediments attributed to the Po river closely correspond to those characteristic of the Po sediments
172 (average 1.81, SD = 0.37) that outcrop southward in the Province of Ferrara (Bianchini et al., 2013).
173 The discriminated series are also characterised by other distinctive siderophile/chalcophile values
174 such as Cr/Pb (Fig. 3b), which is highest in the Po (average 7.95, SD = 2.1), lowest in the Adige
175 (average 1.41, SD = 0.35) and intermediate in the transitional (average 2.39, SD = 0.75) sediments.
176 Three sample groups are also shown in the Ba vs Al₂O₃ diagram (Fig. 3c) that indicates a general
177 enrichment of Ba in the fine fraction, and comparatively higher values in the Adige alluvial
178 sediments with respect to those from the Po sediments. The different concentrations of
179 siderophile/chalcophile elements in the recognised sample groups is generalised in the
180 Cr+Ni+Co+V vs Zn+Pb+Cu+Ba diagram (Fig. 4a) and schematised in the Ni/Zn vs Cr/Pb diagram
181 (Fig. 4b) that provide a geochemical classification scheme for the sediment provenance in the
182 investigated area. The relationships between alkaline and alkaline-earth elements are also used to
183 distinguish the sediments: Po sediments are relatively enriched in Sr (ppm), whereas the Adige
184 sediments are relatively enriched in K₂O (wt%) and Ba (ppm). Further geochemical differences are
185 shown by P₂O₅ and MnO (wt%), which are generally higher in the Adige and Po samples,
186 respectively.

187
188 The bulk rock geochemical composition of 18 bed and suspended sediments from the Po and Adige
189 rivers are presented in Supplementary Table 6 and summarised in Table 2. The major element
190 composition of the bed sediments is, in general, relatively enriched in SiO₂ compared to the
191 suspended sediments, which are enriched in Al₂O₃ and K₂O (wt%). This observation reflects the
192 prevalence of quartz in the coarser bed sediments and of clay minerals and carbonates in the finer
193 suspended particles. The bed sediments are generally depleted in Ni, V, Zn, Pb, and Cu compared to
194 the suspended sediments, which indicates that most heavy metals are preferentially partitioned in
195 clay minerals and/or in the associated amorphous matter. A comparison between sediments from the
196 Po and Adige rivers shows that the former is relatively enriched in CaO compared to MgO (wt%).
197 For the heavy metals, the Po River sediments are characterised by higher concentrations of
198 siderophile elements whereas the Adige River sediments are comparatively enriched in chalcophile
199 elements with geochemical trends overlapping those of their alluvial counterparts (Figs. 3 and 4).
200 Coherently, the Po River sediments have higher Ni/Zn (average 1.46, SD = 0.28) and Cr/Pb
201 (average 5.36, SD = 1.10) ratios compared to the Adige River sediments (average Ni/Zn = 0.32, SD
202 = 0.04; average Cr/Pb = 1.21, SD = 0.38). The Po sediments have a relatively high Sr content
203 (ppm), whereas the Adige sediments are characterised by a comparative Ba enrichment (ppm).

204

205 **Discussion**

206 The reported geochemical data for alluvial sediments from the northeastern Padanian plain provide
207 constraints for the Holocene palaeohydrographic evolution of the area, which is primarily
208 determined by geomorphological studies (often based on remote sensing data), archaeological
209 discoveries and historical chronicles (e.g., Castiglioni, 1999).

210 The major element composition primarily reflects the mineralogical features of distinct depositional
211 facies. The positive correlation between sand content (%) and SiO₂ (wt%) suggests that the coarse
212 fraction (channel deposits) is generally dominated by quartz and feldspars, whereas the increase of
213 clay and SOM contents in parallel with Fe₂O₃, K₂O and Al₂O₃ (wt%) suggests that the fine fraction
214 (distal levee deposits) includes iron-rich clay minerals, such as chlorite and smectite, and that it is
215 intimately associated with the organic matter. However, major elements are scarcely discriminant
216 for the sediment provenance in the study area, where the Po and Adige river systems variably
217 migrated and interfingered during the last millennia. Among the major elements, only CaO and
218 MgO (wt%) give an approximate indication for the sediment provenance, as the CaO/MgO ratio is
219 generally higher in the Po River deposits.

220 In contrast, for the current case study, we have demonstrated that the trace element signature is a
221 precise tool for the identification of the Po and Adige end-members, which refines the
222 discrimination criteria previously proposed by Amorosi et al. (2008) and Picone et al. (2008).
223 According to our inferences, alluvial sediments with Ni/Zn > 1.0 and Cr/Pb > 4.2 are clearly
224 attributable to the Po River basin. This geochemical fingerprint, characterised by a relative
225 enrichment in siderophile elements, is related to the presence of mafic and ultramafic rocks
226 (primarily included in ophiolite sequences) that outcrop within the Po River hydrological basin
227 (Amorosi, 2012; Bianchini et al., 2012; 2013). Conversely, alluvial sediments with Ni/Zn < 0.6 and
228 Cr/Pb < 1.9 are clearly attributable to the Adige River basin, where the relative enrichment of
229 chalcophile elements is related to the presence of Zn-Pb-Cu sulphide ore deposits, which are
230 occasionally associated with barite in the mountainous part of the catchment. These discrimination
231 criteria have been validated by the Po and Adige river sediment data retrieved from the available
232 literature (Boldrin et al., 1989; Dinelli and Lucchini, 1999). The Po soils and sediments are
233 relatively enriched in Sr (ppm) resulting from the weathering of plagioclase-bearing mafic rocks in
234 addition to the geochemical budget related to carbonates. Conversely, Ba (ppm) enrichment of the
235 Adige sediments is possibly related to the weathering of alkaline feldspar-bearing felsic rocks
236 and/or to barite deposits. Moreover, the Adige sediments are generally enriched in P₂O₅ (wt%)
237 whereas the Po sediments are enriched in MnO (wt%). However, the distribution of other elements,
238 such as Zr and REEs, do not differentiate the sediments of the two riverine systems.

239 Samples having transitional characteristics between the Po and Adige geochemical affinities
240 represent a mixture of the above described end-members. They could result from the juxtaposition
241 of several alluvial events for the two fluvial vectors occurring in historical times when both the Po
242 and Adige rivers were scarcely channelled, thus more frequently subjected to overflow and avulsion
243 (Stefani and Vincenzi, 2005). However, remarkable inter-element correlations (Figs. 3 and 4)
244 suggest that the transitional series represents a precise mixture of the Po and Adige sediments in the
245 proportions of 70% and 30%, respectively. This homogeneity implies the existence of an ancient
246 fluvial system collecting water and sediments from both river basins and flowing between the

247 current stream paths of the Po and Adige rivers, which is possibly represented by a series of
248 palaeochannels known as “Po di Fratta”, “Cona-Saline” (Bronze age) and “Po di Adria” (Iron age-
249 Roman times; Piovan et al., 2010; 2012; Ravazzi et al., 2013). The existence of these ancient
250 branches of the Po River has been confirmed by petrographic studies on coarse-grained sediments
251 from alluvial ridges near Rovigo that highlighted the presence of key minerals (glaucofane and
252 serpentine) attributable to the weathering of rocks (blueschists and serpentinites) that outcrop
253 exclusively in the Po River basin. Unfortunately, the approach adopted by previous studies is
254 appropriate only for palaeochannel deposits, whereas it cannot be applied to finer sediments, such
255 as those of distal levee facies, which can only be interpreted using bulk sample geochemistry. Our
256 new data suggests that this ancient Po River stream path possibly captured a branch of the Adige
257 fluvial system (also including minor distributaries oriented southward, Castiglioni, 1999), thus
258 acquiring part of its sedimentary load that effectively mixed and generated the transitional
259 sediments widespread in the study area. However, this hypothesis is scarcely constrained by the
260 available geomorphological evidence (Piovan et al., 2012; Fontana et al., 2014) and historical
261 chronicles (Calzolari, 2008) and, therefore, it is more plausible that the interaction between the Po
262 and Adige fluvial systems (that variably interfingered in space and time during the last millennia)
263 occurred through the development of wetlands fed by both the Po and Adige rivers, especially in
264 Pre-Roman times (Stefani and Vincenzi, 2005). This interpretation is supported by the prevalent
265 fine texture of samples belonging to the transitional series, as well as by historical chronicles that
266 describe the widespread presence of flooded zones in the studied area (Whitehouse, 2001).

267

268 **Conclusions**

269 The geochemical fingerprint of the Province of Rovigo sediments clearly characterises the solid
270 load transported by the two primary riverine systems that envelope the area, i.e., the Po River
271 containing sediments characterised by notable concentrations of Ni and Cr, and the Adige River
272 containing sediments relatively enriched in chalcophile elements such as Pb and Zn. Therefore,
273 geochemistry represents a fundamental tool to correlate the investigated alluvial soils and sediments
274 to a specific basin characterised by distinct parent rocks. Geochemical maps showing the spatial
275 distribution of key elements are useful to define geochemical gradients and local anomalies, but the
276 precise identification of the sediment provenance can only be assessed using key elemental ratios.
277 The current study shows the potential of the geochemical approach in the reconstruction of
278 palaeohydrographic patterns. We demonstrated that the Po River sedimentary influence extends
279 northward compared to its current stream path (as also proposed by authors that studied sediments
280 in the Adriatic Sea; Amorosi et al., 2008; Picone et al. 2008) and that the Po River sediments are
281 locally present very close to the modern Adige stream path. Moreover, the presented geochemical
282 data suggest that in historical times a branch of the Po River crosscutting the studied area was
283 characterised by a sedimentary load with a hybrid signature intermediate between those currently
284 recognised in the Po and Adige rivers. These transitional sediments may suggest that at that time the
285 two fluvial systems were connected, i.e., that a branch of the Adige was a tributary of the Po River
286 upstream with respect to the study area or, more probably, that the interference of the two fluvial
287 systems developed through marsh/swamp environments that are recorded in historical chronicles
288 and in geomorphological reconstructions (Vincenzi and Stefani, 2005). These sedimentary facies

289 could, therefore, represent the expression of depositional domains in which both the Po and Adige
290 rivers delivered sedimentary inputs that probably originated within wetlands developed among
291 major fluvial ridges.

292

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296

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298

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404

405 **Table captions:**

406

407 **Table 1** – Geochemical composition of soils and sediments from the investigated sector of the
408 Padanian plain (Province of Rovigo, NE Italy).

409

410 **Table 2** – Average geochemical composition of modern bed and suspended sediments from the Po
411 and Adige rivers.

412

413

414

Research Highlights:

- Padanian Plain landscape results from the interplay of complex fluvial systems
- Po and Adige river systems variously interacted in space and time in the last millennia
- Geochemical analyses allow discrimination between Po and Adige river sediments
- Ni/Zn and Cr/Pb are useful geochemical proxies of sediment provenance
- Geochemical proxies are useful for palaeohydrographic reconstructions

Table 1
[Click here to download Table: Table 1.docx](#)

	Boara Polesine				Concadirame				Lusia				Fratra Polesine-Bosaro				Stienta			
	Median	Mean	1 st Quart	3 rd Quart	Median	Mean	1 st Quart	3 rd Quart	Median	Mean	1 st Quart	3 rd Quart	Median	Mean	1 st Quart	3 rd Quart	Median	Mean	1 st Quart	3 rd Quart
SiO₂ (wt%)	49.15	49.17	47.12	51.60	51.96	51.97	49.85	54.02	51.22	53.61	48.82	61.34	52.63	52.60	51.94	53.66	51.99	53.65	50.38	55.76
TiO₂	0.69	0.67	0.67	0.71	0.67	0.66	0.63	0.69	0.65	0.62	0.52	0.68	0.63	0.63	0.60	0.65	0.61	0.59	0.58	0.63
Al₂O₃	15.43	14.82	13.46	16.28	13.31	13.28	12.67	14.01	13.32	12.57	10.33	13.90	13.52	13.67	13.30	13.91	12.88	12.68	12.12	13.30
Fe₂O₃	7.36	7.33	5.88	8.65	5.50	5.65	5.16	6.09	5.57	5.15	4.06	5.79	4.92	4.87	4.61	5.16	5.91	5.85	5.40	6.34
MnO	0.09	0.09	0.08	0.10	0.09	0.09	0.09	0.10	0.09	0.08	0.07	0.09	0.09	0.09	0.09	0.10	0.12	0.11	0.10	0.12
MgO	4.47	4.20	3.74	4.76	4.97	4.93	4.58	5.20	5.25	4.87	3.93	5.46	4.73	4.76	4.53	4.90	4.58	4.45	4.34	4.78
CaO	6.35	5.75	3.03	7.65	8.04	7.98	6.97	9.09	7.72	8.17	7.46	8.26	8.82	8.48	8.07	8.97	8.55	8.56	7.92	9.38
Na₂O	0.82	0.89	0.72	1.00	1.12	1.11	1.02	1.20	1.12	1.23	1.05	1.52	1.11	1.14	1.02	1.20	0.87	0.93	0.76	1.04
K₂O	3.21	3.12	3.05	3.33	2.88	2.87	2.77	3.02	2.98	2.94	2.76	3.10	2.40	2.49	2.27	2.64	2.41	2.37	2.20	2.53
P₂O₅	0.19	0.20	0.16	0.22	0.21	0.23	0.17	0.28	0.28	0.28	0.21	0.36	0.23	0.25	0.18	0.28	0.14	0.14	0.11	0.16
LOI	12.46	13.75	11.12	15.43	11.46	11.24	9.66	12.17	11.21	10.47	7.85	12.31	11.01	11.02	10.45	11.57	11.21	10.66	9.66	12.34
Ba (ppm)	512	497	469	554	471	466	448	498	497	460	348	529	400	417	380	440	358	359	337	382
Ce	57	60	51	71	45	46	43	50	49	45	39	51	50	51	47	53	41	42	40	45
Co	24	24	19	29	17	18	15	20	18	17	12	20	19	19	16	22	25	24	22	26
Cr	129	128	83	154	75	93	65	95	73	69	56	80	178	157	108	197	200	192	178	206
Cu	59	64	46	86	38	37	29	45	40	50	34	52	40	44	32	48	46	46	32	51
Ga	24	24	20	28	18	18	16	19	20	18	14	20	17	17	16	18	16	15	14	18
Hf	8	8	7	9	8	8	7	9	8	8	7	8	4	4	4	5	4	4	4	5
La	38	36	27	41	27	27	23	30	25	24	20	27	28	29	24	30	27	26	23	31
Nb	15	15	14	16	15	15	14	16	14	14	12	15	14	15	13	16	10	10	9	13
Nd	32	30	25	37	22	23	20	26	21	21	19	23	27	28	25	30	20	20	18	22
Ni	98	94	64	119	42	57	36	53	39	38	21	48	133	118	76	154	143	144	130	166
Pb	53	57	40	73	49	51	45	57	47	50	33	64	28	36	23	38	24	26	21	29
Rb	143	146	111	184	106	108	102	113	113	109	96	119	106	106	99	114	88	84	73	97
Sc	18	18	16	20	15	15	14	17	15	14	11	16	12	13	12	13	17	17	16	18
Sr	181	181	168	200	189	192	182	198	192	189	168	202	209	213	190	242	250	240	197	296
Th	11	11	9	13	8	8	8	9	8	8	6	9	7	7	6	8	7	7	7	8
V	129	123	96	146	89	89	79	99	96	83	47	102	93	92	84	100	101	99	92	110
Y	24	24	22	27	23	23	20	26	24	22	16	25	26	26	24	28	19	18	15	22
Zn	131	137	108	159	96	105	86	117	97	98	73	115	88	92	78	99	84	90	70	94
Zr	127	132	109	159	169	169	152	182	173	158	127	187	263	268	249	302	130	122	103	147
SOM (wt%)	8.5	9.7	5.9	12.4	11.5	11.2	9.7	12.2	-	-	-	-	3.2	3.0	2.5	3.7	4.6	4.4	3.2	6.0

Footnote: LOI = Loss On Ignition, SOM = Soil Organic Matter, - = not determined.

Table 2

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	Po river		Adige river	
	<i>Bed sediments</i>	<i>Suspended sediments</i>	<i>Bed sediments</i>	<i>Suspended sediments</i>
SiO₂ (wt%)	52.74	47.50	53.46	37.16
TiO₂	0.62	0.62	0.60	0.49
Al₂O₃	12.10	14.00	11.94	13.02
Fe₂O₃	4.54	5.46	3.73	4.93
MnO	0.13	0.12	0.07	0.09
MgO	5.86	4.57	6.31	8.77
CaO	10.53	9.80	9.23	12.27
Na₂O	1.20	0.80	1.45	1.07
K₂O	2.08	2.59	2.54	2.62
P₂O₅	0.21	0.18	0.22	0.17
LOI	10.01	14.37	10.47	19.43
Ba (ppm)	360	424	481	546
Ce	37	52	41	56
Co	27	23	13	16
Cr	241	197	53	70
Cu	33	57	20	57
Ga	14	22	15	21
Hf	5	4	4	2
La	24	33	23	37
Nb	11	15	12	10
Nd	25	21	27	21
Ni	156	187	31	60
Pb	28	36	37	68
Rb	87	129	102	112
Sc	9	14	9	9
Sr	291	321	186	198
Th	6	7	5	3
V	77	117	66	99
Y	25	24	24	23
Zn	97	124	93	202
Zr	227	197	250	180

Footnote: LOI = Loss On Ignition.

Figure 1

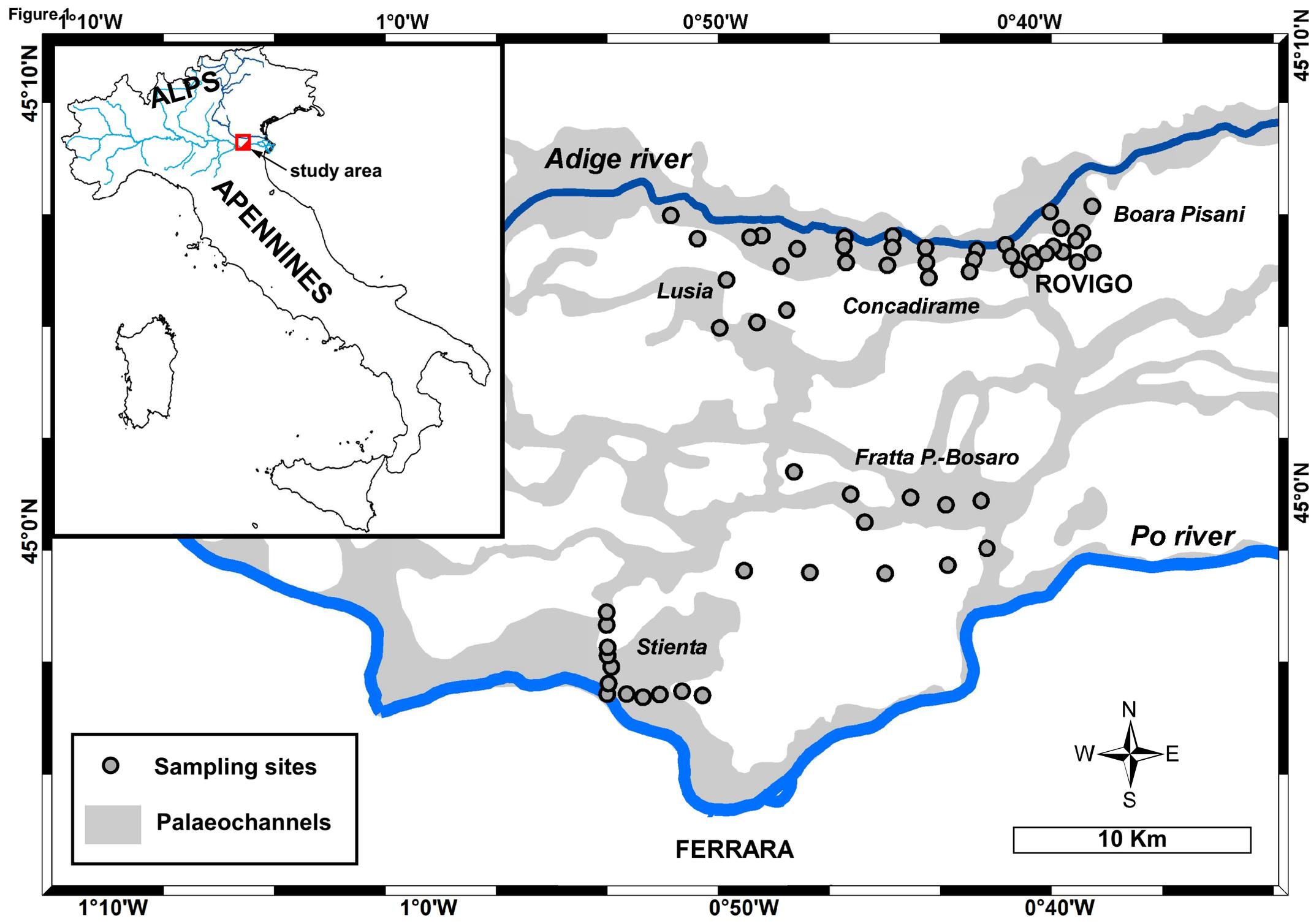


Figure 2

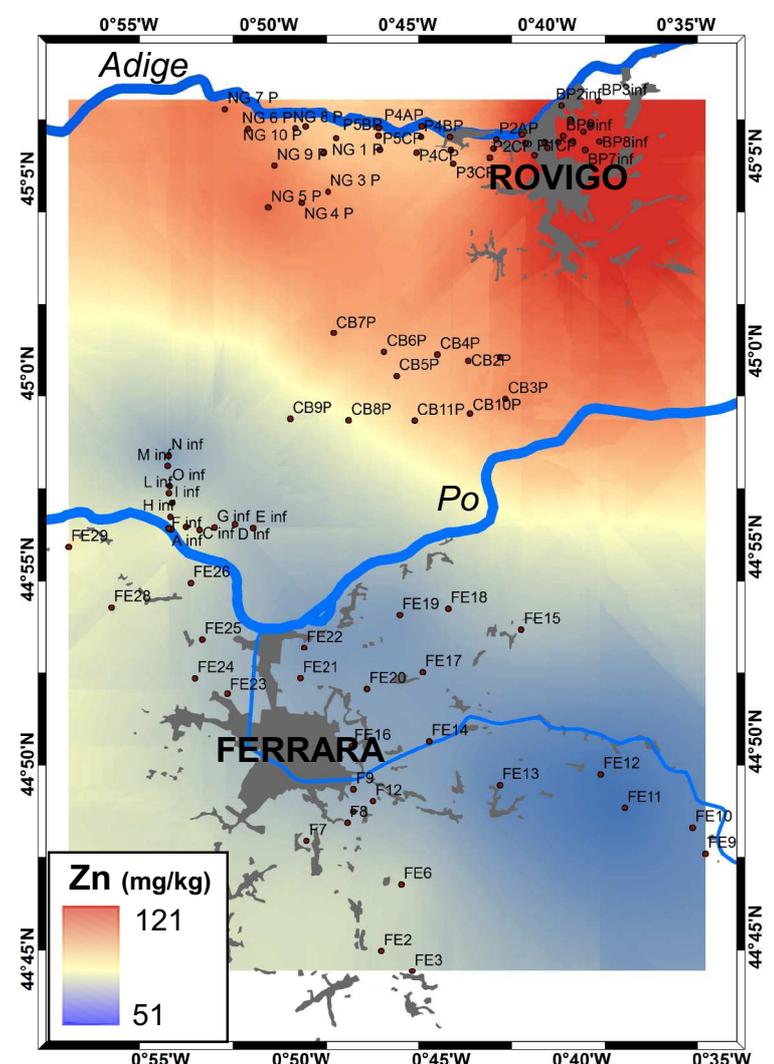
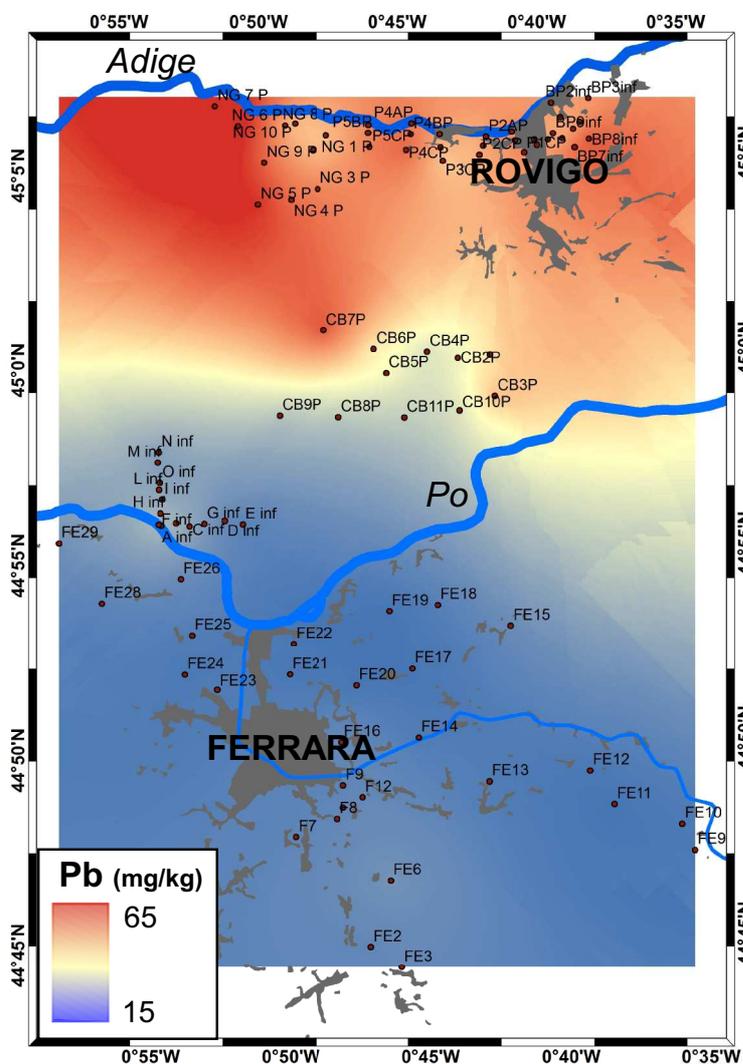
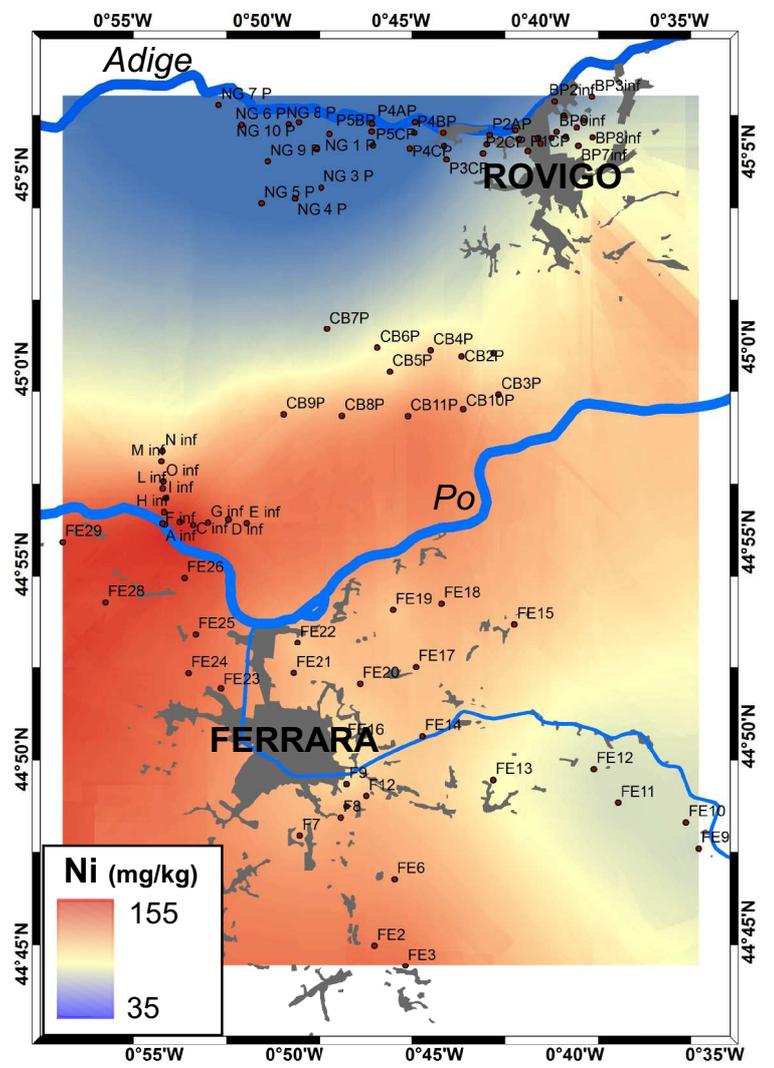
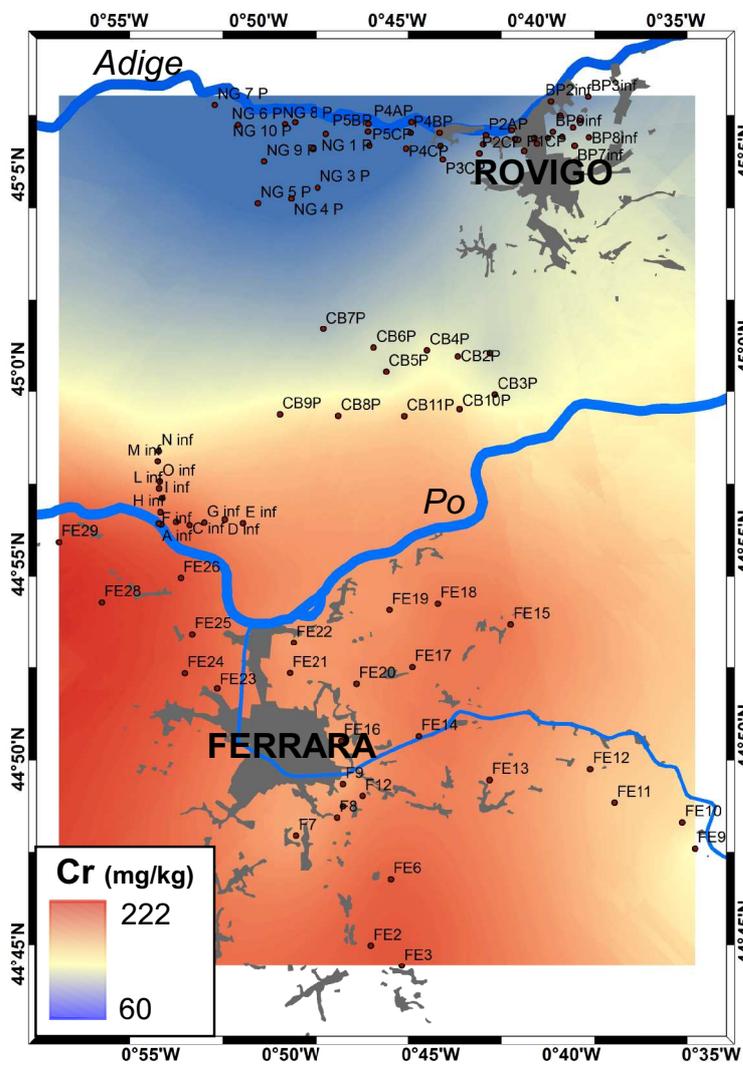


Figure 3 *Po river basin*

- alluvial sediments
- ◆ bed river sediments
- suspended river sediments

Adige river basin

- alluvial sediments
- ◇ bed river sediments
- suspended river sediments

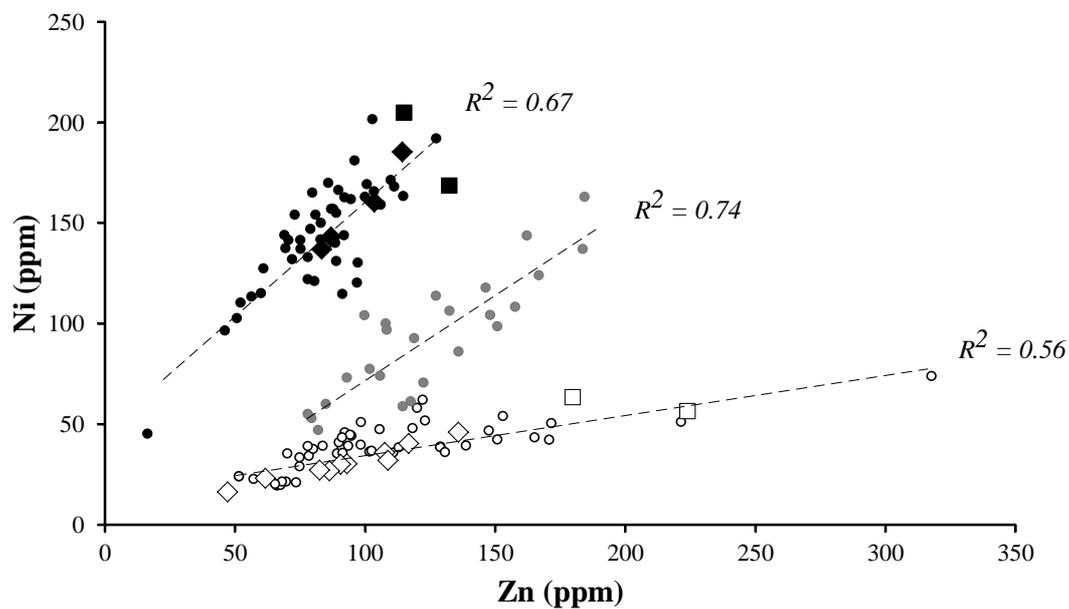
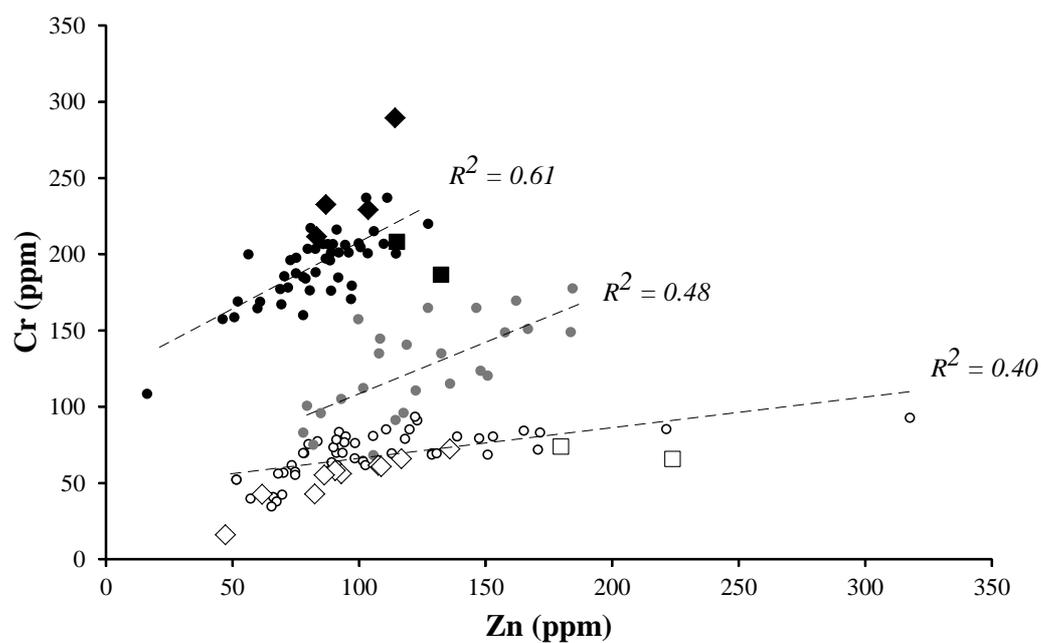
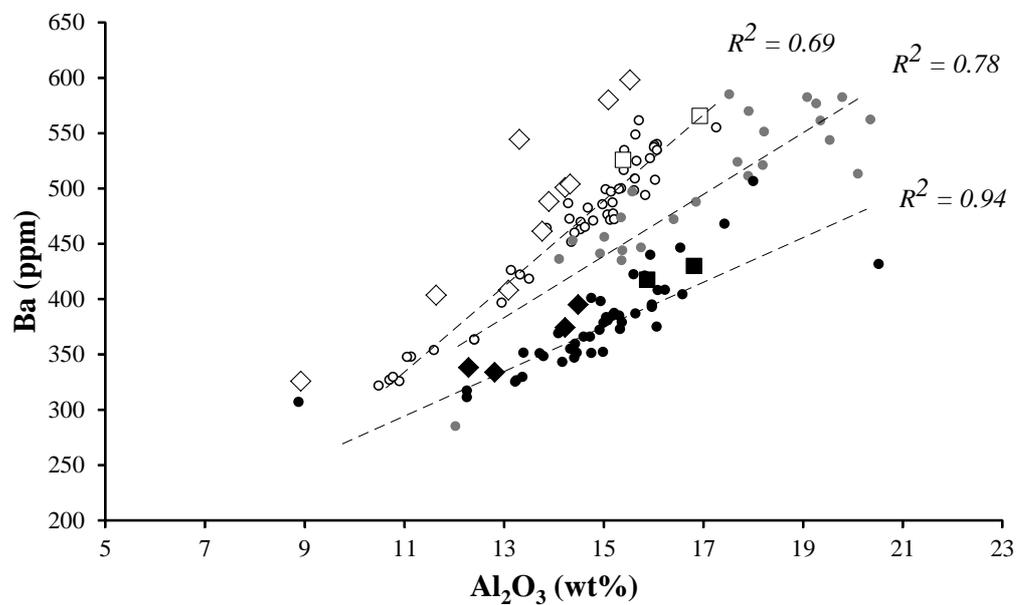
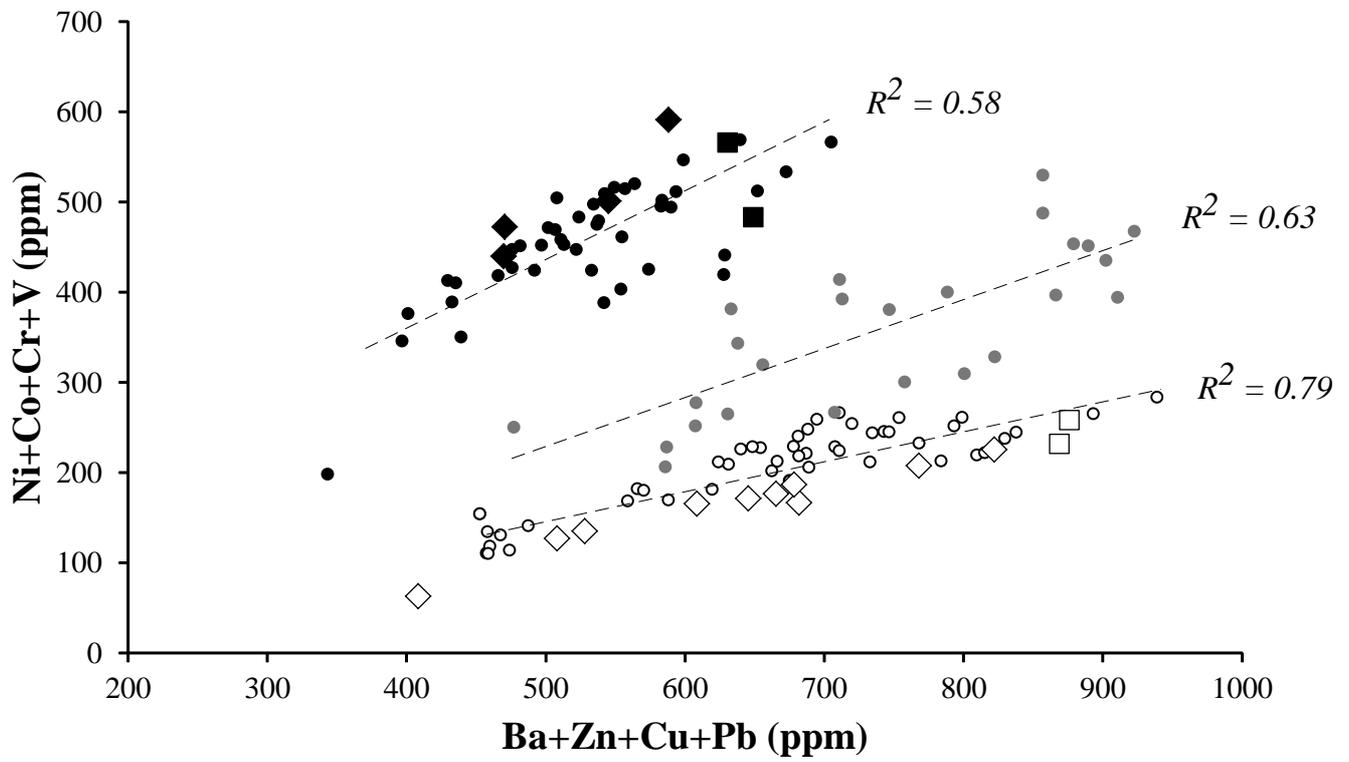
a)**b)****c)**

Figure 4

a)



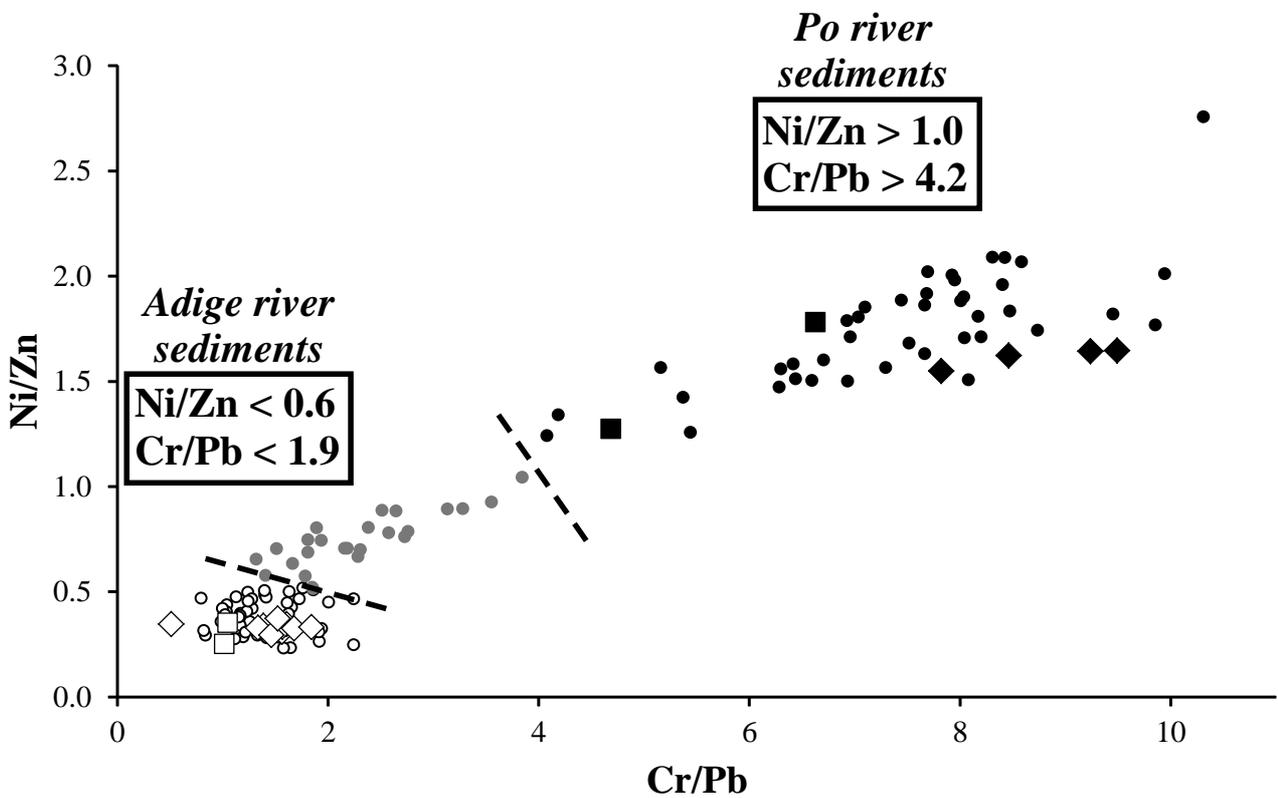
Po river basin

- alluvial sediments
- ◆ bed river sediments
- suspended river sediments

Adige river basin

- alluvial sediments
- ◇ bed river sediments
- suspended river sediments

b)



1 **Figure captions:**

2
3 **Figure 1** – Sketch map of the studied area reporting the Po and Adige river paths, sampling
4 locations of the studied alluvial sediments and palaeochannel distribution retrieved from the Veneto
5 regional soil map 1:250000 (Osservatorio Regionale Suolo dell'ARPAV, 2015). Gauss-Boaga
6 projection (W), datum: Rome 1940.

7
8 **Figure 2** – Geochemical maps highlighting the spatial distribution of Cr, Ni, Pb and Zn (ppm) in
9 alluvial sediments (collected at 90-110 cm depth) from the studied sector of the Padanian plain.
10 Data from this work and from Bianchini et al. (2013). See text for further details.

11
12 **Figure 3** – Ni vs Zn (a), Cr vs Pb (b) and Ba vs Al₂O₃ (c) bivariate diagrams highlighting the
13 geochemical features of Po and Adige river sediments. The diagrams report composition of alluvial
14 sediments from the study area (see Supplementary Tables 1-5) and modern bed and suspended
15 sediment from Adige and Po rivers (Supplementary Table 6). Note that the distribution of alluvial
16 sediments identifies three distinct geochemical trends, the most extreme of them conforming to Po
17 and Adige river sediments. The intermediate trend can be interpreted as a mixture of 70% Po and
18 30% Adige sedimentary end-members.

19
20 **Figure 4** – Discrimination diagrams based on (a) the sum of siderophile (Cr+Ni+Co+V) vs
21 chalcophile (Zn+Cu+Pb+Ba) elements and (b) ratios between siderophile (Ni, Cr) vs chalcophile
22 (Zn, Pb) elements.

23

Supplementary Tables 1-6

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