




## Research Article

# Tropical island adaptations in Southeast Asia during the Last Glacial Maximum: evidence from Palawan

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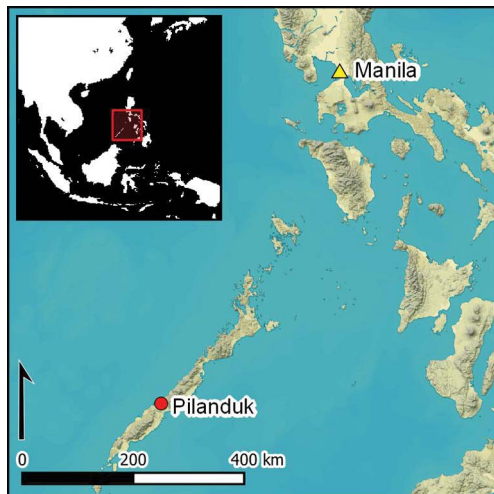
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Through the Late Pleistocene and Early Holocene, humans have adapted to significant climate and environmental change. One key region for investigating these adaptive strategies is Island Southeast Asia, where fluctuating sea levels led to dramatic changes in coastlines, vegetation and fauna. The authors present new radiocarbon dates and faunal data from the re-excavation of Pilanduk Cave on Palawan Island, Philippines. The results corroborate the results of earlier excavations that identified Pleistocene occupation of the site. Pilanduk shows evidence for specialised deer hunting and freshwater mollusc consumption during the Last Glacial Maximum. The results add to the evidence for the shifting foraging behaviours of modern humans occupying variable tropical environments across Island Southeast Asia.

Keywords: Philippines, Pilanduk Cave, Palaeolithic, MIS 2, zooarchaeology, subsistence, foragers

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## Introduction

Island Southeast Asia has emerged as a key region for the investigation of hominin evolutionary trends and human adaptations to tropical environments, including rainforest ecosystems and maritime island settings (O'Connor *et al.* 2017; Roberts & Stewart 2018). Early records of modern human settlement in Southeast Asia show evidence for the occupation of tropical rainforests in Borneo and Sumatra (Barker *et al.* 2017; Westaway *et al.* 2017) during Marine Isotope Stages (MIS) 3 and 4, respectively (*c.* 71 000–30 000 years ago). During this period,

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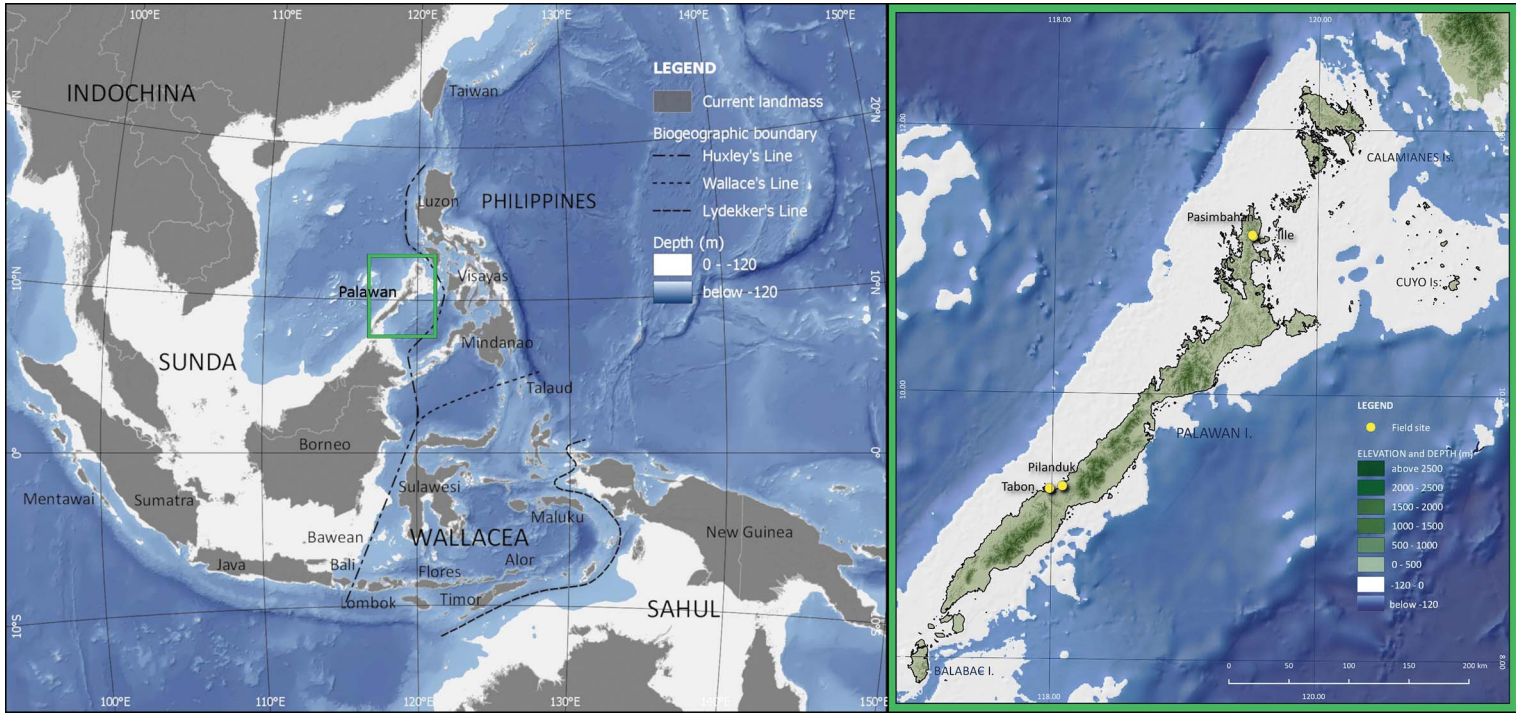
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45 environments across the region changed dramatically, necessitating adaptation by human  
 46 populations. A globally important period of change was during the Last Glacial Maximum  
 47 (LGM; 30 000–19 000 years ago), when certain parts of Southeast Asia experienced increas-  
 48 ing aridity and the expansion of open environments (e.g. Wurster *et al.* 2019). A number of  
 49 archaeological sites in the Malay Peninsula and Sulawesi demonstrate abandonment in  
 50 response to this increasing aridity, suggesting that some populations may have relocated to  
 51 coastal areas (O’Connor & Bulbeck 2014). Many of these LGM coastlines and locales across  
 52 Southeast Asia, which were exposed during the sea-level lowstand, are currently submerged  
 53 (Sathiamurthy & Voris 2006). This partially explains why regional archaeological data for  
 54 the LGM are relatively scarce and less known compared to Terminal Pleistocene and Holo-  
 55 cene sequences (O’Connor & Bulbeck 2014). For this reason, a rare and well-preserved Phil-  
 56ippine LGM record affords a valuable opportunity to investigate human responses to  
 57 extensive palaeoenvironmental changes in tropical ecosystems. In this article, we examine  
 58 these adaptations at the site of Pilanduk Cave, Palawan Island (Figure 1). We then synthesise  
 59 the data from this site in the context of the wider Palawan subsistence record in order to ana-  
 60lyse the cultural and adaptive plasticity that has been attributed to Late Pleistocene *Homo*  
 61 *sapiens* (Roberts & Amano 2019). The extensive and flexible behavioural repertoire of mod-  
 62ern humans manifests in the colonisation of various environments and habitats, including the  
 63 tropical and maritime ecosystems of Island Southeast Asia. The study of these changing eco-  
 64 systems provides temporal depth to the investigation of how human adaptations varied in the  
 65 tropics.

66 The LGM was a period of drastic climatic and environmental change across the globe  
 67 (Lambeck *et al.* 2014). In Southeast Asia, the Sunda Shelf was exposed and Sundaland  
 68 reached its maximal extent during the LGM, when it is estimated that sea levels were as  
 69 low as  $-123 \pm 2$  m bpsl (metres below present-day sea level; Figure 1) (Sathiamurthy &  
 70 Voris 2006; Hanebuth *et al.* 2009). With the exposure of this continental shelf, a reorgan-  
 71 isation of oceanic currents and atmospheric systems took effect across the Indo-Pacific Warm  
 72 Pool, and sea surface temperatures decreased by 2–3°C compared to the present day (De  
 73 Deckker *et al.* 2003; Gagan *et al.* 2004). A ‘savanna corridor’ hypothesis has long been pro-  
 74 posed for central Sundaland and the western Philippines (Heaney 1991), which states, that  
 75 during the LGM, a wide expanse of savanna extended down the Malaysian Peninsula and  
 76 across the submerged landmass between Borneo and Java. LGM palaeoenvironmental data  
 77 appear to support such a scenario (Bird *et al.* 2005; Wurster *et al.* 2019), although other scho-  
 78 lars suggest the continued persistence of closed tropical rainforests in parts of the exposed  
 79 Sunda Shelf (Wang *et al.* 2009). Other areas in Southeast Asia retained stretches of lowland  
 80 tropical forests during the LGM, such as at the Niah Caves of Borneo (Barker *et al.* 2017) and  
 81 Trang An in northern Vietnam (Rabett *et al.* 2017).

82 On Palawan, palaeoenvironmental records document drier conditions during the LGM,  
 83 which is bracketed by wetter and more humid conditions characterised by closed forest dur-  
 84 ing MIS 3 and the Holocene. These conditions are attested by stable isotope analysis of  
 85 ancient guano deposits from three caves across the island. In Tabon Cave, analysis of three  
 86 Pleistocene phases of guano deposits reveals environments that were predominantly charac-  
 87 terised by closed forest during MIS 3 (Choa 2018). The youngest of these sequences corre-  
 88 sponds in age with a Tabon archaeological layer dated to *c.* 39 000 cal BP. In Makangit Cave

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**Q5** Figure 1. Map of Southeast Asia (left) and Palawan Island (right) showing present-day land distribution, biogeographic regions, biogeographic boundaries and the location of Pilanduk Cave and other Palawan sites mentioned in the text. The  $-120\text{m}$  isobath approximates the possible extent of landmasses during the Last Glacial Maximum, including that of the Greater Palawan landmass (© J. Ochoa and J. Villasper).

133 in northern Palawan and Gangub Cave in the south, stable isotope data signal a much drier  
 134 climate and savanna conditions, dominated by  $C_4$  vegetation, during the LGM (Bird *et al.*  
 135 2007; Wurster *et al.* 2010). These data also indicate that closed rainforest expanded again  
 136 *c.* 13 500 cal BP and that, by the Early Holocene, the savanna vegetation was replaced by  
 137 closed tropical forest—similar to that of the present day. Palaeogeographic reconstructions  
 138 also indicate that Palawan was connected by land to adjacent, smaller islands (Balabac,  
 139 Cuyo and Calamianes) during the LGM, forming the landmass known as ‘Greater Palawan’  
 140 (Figure 1). A narrow land-bridge connection with Borneo may have been possible with a sea-  
 141 level drop of at least  $-135$  m (Robles *et al.* 2015); estimates by Lambeck *et al.* (2014) indicate  
 142 a sea-level lowstand of approximately  $-134$ m at *c.* 21 000 cal BP.

## 144 Pilanduk Cave and Palawan Island archaeology

145  
 146 Palawan is the fifth largest island in the Philippine archipelago. It lies on the north-eastern tip  
 147 of the Sunda Shelf, west of Huxley’s Line, the biogeographic boundary demarcating Sunda  
 148 from Wallacea (Figure 1). The oldest known human fossils on Palawan are from Tabon Cave  
 149 and are ascribed to *Homo sapiens* (Fox 1970; Détroit *et al.* 2004). Direct dating of these  
 150 human remains has yielded ages ranging from  $16\ 500 \pm 2000$  to  $47\ 000 \pm 11\ 000$  years BP  
 151 (Détroit *et al.* 2004). Additional radiometric dates obtained by Choa (2018) for the  
 152 Tabon stratigraphy have yielded ages of *c.* 39 000–33 000 cal BP for what Fox (1970) cate-  
 153 gorised as Flake Assemblages II and III. Palawan is currently the only island in the Philippine  
 154 archipelago with Pleistocene *H. sapiens* fossils. In northern Palawan, Ille Cave and Pasimba-  
 155 han Cave have produced archaeological sequences spanning the last 14 000 years and 10 000  
 156 years, respectively (Lewis *et al.* 2008; Ochoa *et al.* 2014).

157 Pilanduk Cave is one of a few Philippine sites that has produced a large and well-preserved  
 158 archaeological assemblage of Late Pleistocene date. The cave is located in Negmisi Peak  
 159 (Magmisi or Devel Peak) at 165m asl and situated near the Iwahig River (Figure 2). Negmisi  
 160 and Pilanduk are within the scope of the ancestral domain of the Indigenous Pala’wan  
 161 (Certificate of Ancestral Domain Title No. RO4-QUE-O110-143) in Barangay Maasin of  
 162 Quezon Municipality. The site is currently less than 1km from the coast (Figure 2c). *Pilan-*  
 163 *duk* is a vernacular term for the mouse-deer (*Tragulus nigricans*) but is also sometimes used to  
 164 refer to the Calamian hog deer (*Axis calamianensis*). The first archaeological excavation at the  
 165 cave was conducted in 1969–1970 by a team led by Jonathan Kress. Three radiocarbon dates  
 166 taken from shell samples (of unreported taxon) are provided by Kress (2000). The deepest  
 167 layer, Layer IV, is dated at  $25\ 470 \pm 1000$  BP. The younger layers, Layers II and I, date to  
 168  $18\ 340 \pm 370$  BP and  $18\ 260 \pm 650$  BP, respectively. Kress (1978) reports that, after the  
 169 Late Pleistocene occupation, the site was rarely used again until the Late Holocene, when  
 170 it was used as a jar-burial site. During its LGM occupation, the site would have been further  
 171 inland, at approximately 40km from the coast, due to the MIS-2 sea-level lowstand (Robles  
 172 *et al.* 2015) (Figure 1).

173 The lithics, pottery and shells recovered in the 1970 excavation were described by Kress in  
 174 three separate publications (Kress 1977, 1978, 2000). Regarding the vertebrate assemblage,  
 175 there are no published faunal counts, although Kress (1977) noted that deer bones dominate  
 176 the Palaeolithic layers. For the mollusc assemblage, Kress (2000) identified 31 species, with

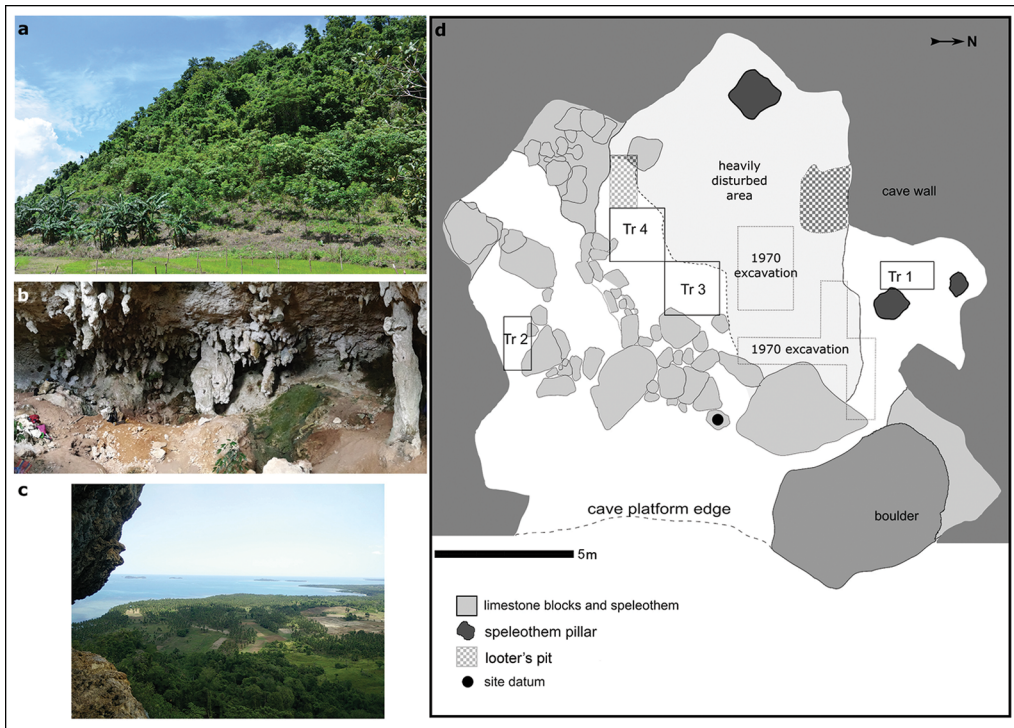


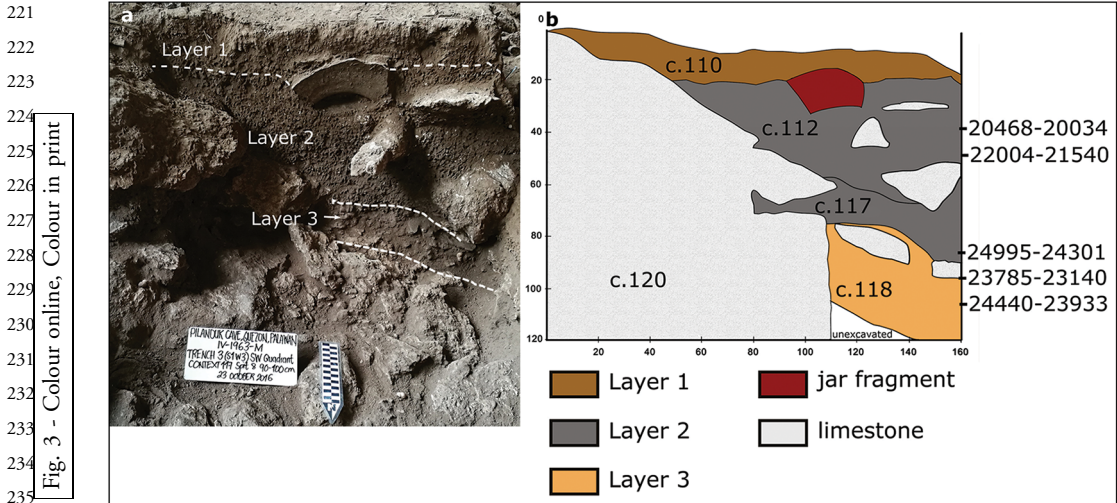
Fig. 2 - Colour online, Colour in print

Figure 2. a) Negmisi Peak viewed from the east by the village road; b) Pilanduk Cave as viewed from the north-east; c) view near the peak of Negmisi showing present-day near-coast environment; d) plan of the 2016 excavation, showing the four trenches and reconstructed location of the 1970 trenches from Kress (2000) (photographs by J. Ochoa; plan drawn by M. Lara, A. Peñalosa and J. Ochoa).

90 per cent being freshwater and terrestrial taxa. This pattern is taken to corroborate the inland environment of the Palaeolithic levels of the sequence. There has been a need to verify the dates reported by Kress, however, due to the limited stratigraphic data available for Pilanduk, and the limitations of the radiocarbon dating method at the time of Kress' excavation in the 1970s, particularly as applied to molluscs.

**Re-excavation and chronology of Pilanduk Cave**

In order to validate the Late Pleistocene age of the site, in October 2016, a re-excavation of Pilanduk Cave was undertaken by an archaeological team composed of members from the National Museum of the Philippines, the University of the Philippines and the local Pala'wan of Barangay Maasin, Quezon Municipality. Four excavation units were opened (Trenches 1–4) (Figure 2). The online supplementary material (OSM) provides details of the excavation and zooarchaeological methods. Trench 3 reached a depth of 1.2m and provided the reconstructed stratigraphic sequence for the site (Figure 3). We identified three main archaeological layers: a Late Holocene layer and two distinct LGM midden layers. The uppermost layer (Layer 1; includes context 110 in Trench 3) contained an earthenware ceramic assemblage and human remains, which also cut into underlying Pleistocene midden deposits (context 111);



236 *Figure 3. Photograph (a) and drawing (b) of south wall of Trench 3, showing the three major layers, corresponding*  
 237 *context numbers (e.g. c.110) and stratigraphic position of radiocarbon dates (cal BP). Scale in cm (© J. Ochoa).*  
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240 in Trench 3, the matrix was partly comprised of loose yellowish-brown silt (Figure 3). We  
 241 observed this layer in all four trenches. This uppermost layer was highly disturbed and admixed,  
 242 containing LGM midden deposits, Late Holocene artefacts, human remains that are of a dis-  
 243 tinctly different preservation state from that of the Late Pleistocene bones, and backfill resulting  
 244 from recent looting. The Pilanduk stratigraphy is unusual in that there appears to be little sedi-  
 245 ment accumulation after the LGM occupation, with the LGM deposits visible only a few centi-  
 246 metres beneath the present-day surface level in the cave. The ceramic assemblage comprises at  
 247 least five large ceremonial vessels, three lids and eight small vessels (Figure S1). On the basis of  
 248 the associated presence of fragmented human remains, some of these vessels appear to have been  
 249 used as burial jars, which were interred within the upper LGM midden accumulation. These  
 250 results coincide with the earlier findings of Kress (1978), which showed that jar-burial remains  
 251 were present in the uppermost levels of the site. A total of 126 human bone fragments were  
 252 identified in the uppermost layer. These human remains clearly differ in their state of preser-  
 253 vation from that of the animal bones from the LGM midden, with the latter being mineralised  
 254 and showing varying degrees of heat alteration. Direct dates have not yet been obtained for this  
 255 layer, but ceramic vessels show affinities to jar-burial assemblages found across Palawan, loosely  
 256 designated as ‘Late Neolithic’ and ‘Metal Period’ (c. 2500 to 2000 BP) (Fox 1970) (Figure S1 in  
 257 the OSM).

258 The second archaeological layer (contexts 111, 112 and 117 in Trench 3; see Figure 3)  
 259 consisted of a dense midden deposit dominated by faunal remains, found across the cave;  
 260 in Trenches 3 and 4, the layer comprised a matrix of greyish-brown silt. This layer was  
 261 also observed in Trenches 1 and 2. Layer 2 contained most of the archaeological material  
 262 from the site, including numerous lithics and faunal remains. An initial study of the lithic  
 263 artefacts shows that brown chert was the main raw material used (Table S1; Figure S2).  
 264 The presence of cores, flakes and waste by-products of varying sizes in Layer 2 (Table S1)

Table 1. Radiocarbon dates for Pilanduk Cave. All samples are from Trench 3 wood charcoal specimens. Ages were calibrated using OxCal v4.3 and IntCal13 and are at 95.4% probability (Reimer *et al.* 2013; Bronk Ramsey 2017). LDP = local datum point.

Accession Number	Lab Code (OxA)	Taxon	Context	Depth from LDP (m)	Radiocarbon age (BP)	Calibration age range (cal BP)
5397	37182	<i>Michelia</i> sp.	112	0.46	16 785±65	20 468–20 034
5587	36553	<i>Wrightia</i> sp.	112	0.52	17 980 ± 70	22 004–21 540
6210	36332	<i>Pinus</i> sp.	117	0.90	20 460 ± 90	24 995–24 301
6255	37180	<i>Kibatalia</i> sp.	118	0.98	19 500 ± 90	23 785–23 140
6255	37181	<i>Kibatalia</i> sp.	118	0.98	19 570 ± 80	23 875–23 295
6257	36331	<i>Syzygium</i> sp.	118	1.03	20 120 ±90	24 440–23 933

suggests that all knapping stages are represented. Extensive evidence for burning in this layer includes bone charring and calcination, and the ubiquitous presence of ash deposits and wood charcoal. These deposits potentially represent the remains of ancient hearths. Kress (1977, 2000) also reported the abundance of lithic and faunal material in the LGM deposits excavated in trenches located in the centre of the cave (Figure 2d). These deposits potentially correspond to Layer 2.

The third archaeological layer (context 118) consisted of a yellowish-brown silt matrix with angular limestone fragments and noticeably fewer ash deposits, bones and lithics compared with the overlying layer. Layer 3 was observed only in Trench 3, since this unit reached lower depths compared to the adjacent Trench 4. For Trenches 1 and 2, which are on the opposite sides of the central cave area (Figure 2d), bedrock was reached after Layer 2 and Layer 3 was not observed.

Radiocarbon dates for Layers 2 and 3 were obtained from wood charcoal (Table 1). All dating samples were hand-collected specimens recovered *in situ* (all dates were calibrated using OxCal v4.3.2 and the IntCal13 calibration curve (Reimer *et al.* 2013; Bronk Ramsey 2017)). Two charcoal samples from context 112 (Layer 2) produced dates that range from *c.* 20 000 to 22 000 cal BP. Two samples from context 118 (Layer 3) produced ages within the range of *c.* 23 000–24 500 cal BP. A sample from the base of context 117 produced an age range of 24 995–24 301 cal BP (OxA-36332). The sedimentary composition, lithics and faunal remains of context 117 are very similar to those of contexts 112 and 111, and differ from those of context 118; the real age of the materials from context 117 could therefore be close to that of materials from context 112. The age bracket obtained for context 117, however, overlaps with a date from context 118 (OxA-36331). This suggests that the context 117 sample, which was taken at the base of this deposit, may have migrated or be derived from the underlying layer. Nonetheless, the new dates obtained from the 2016 excavation confirm the LGM dating of the site, as previously reported by Kress (2000).

## Vertebrate assemblage

A total of 8491 terrestrial vertebrate specimens have been analysed from Pilanduk Cave, of which 7826 are from the 2016 re-excavation of the site (Table 2). The remaining 665 fragments come from reanalysis of the 1970 assemblage, although few of these fragments

Taxon	Trench 3 Context				2016 Total	Kress' excavation					1970		% NISP	
	Surface	111	112	117		118	Layer 1	Layer 2	Layer 3	Layer 4	No layer	total		Total
<i>Axis calamianensis</i> *	19	7	30	19	1	76	1				8	9	85	1
<i>Rusa</i> sp.*	80	29	53	13	13	188	10	6		7	40	63	251	3
Cervid*	128	163	239	266	38	834	39	8	37	5	106	195	1029	12.1
<i>Sus ahoenobarbus</i>	5	14	15	28	0	62	8	3	7		11	29	91	1.1
<i>Panthera tigris</i> *	1	1	4	1		9**							9	0.1
large mammal	28	979	1004	1718	266	3995	65	8	163	8	59	303	4298	50.6
<i>Macaca fascicularis</i>	1			1		2					2	2	4	0.05
Sciuridae				1	1	2							2	0.02
<i>Hystrix pumila</i>			2			2							2	0.02
intermediate mammal		3	1	5	3	12	1		1			2	14	0.2
<i>Hipposideros diadema</i>					4	4							4	0.05
Chiroptera					22	22							22	0.3
<i>Varanus</i> cf. <i>palawanensis</i>				2		2							2	0.02
Geoemydid		1	8	20	5	34			2		1	3	37	0.4
<i>Cyclemys dentata</i>			3	2		5							5	0.1
Bird		3		3	17	23							23	0.3
Macrovertebrate		971	962	483	106	2522	3		55		1	59	2581	30.4
Microvertebrate		2			30	32							32	0.4
	261	2173	2321	2562	506	7826	127	25	265	20	228	665	8491	100.0

\*Locally extinct taxa.

\*\*Two tiger specimens were from Trench 4.



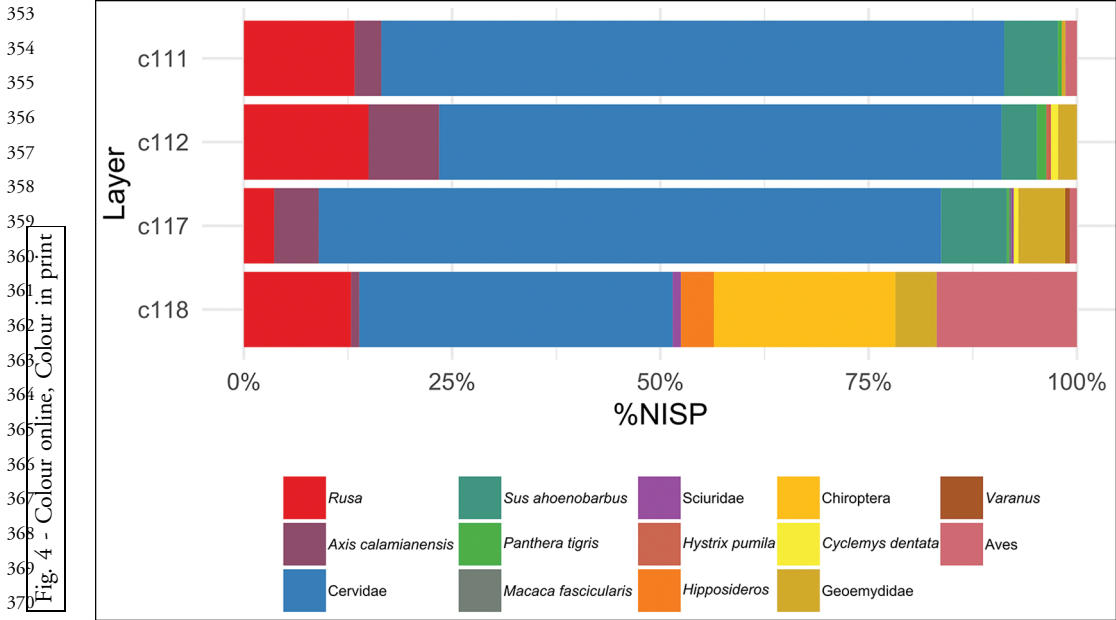


Figure 4. Relative taxonomic abundance (%NISP) of vertebrate taxa at Pilanduk Cave across four major archaeological contexts (c). NISP = number of identified specimens. NISP data are in Table 2 (© J. Ochoa).

transpired to have sufficient stratigraphic information for temporal analysis. Analysis of the 2016 assemblage focused on the Trench 3 stratigraphic sequence, as this unit had the longest sequence of all the trenches. Seven mammalian taxa and two reptile taxa were identified (Table 2). The Pilanduk evidence currently represents the oldest known fossil records of these nine taxa. Cervid remains dominate the LGM assemblage (Figure 4). Two species are represented: a sambar-type deer (*Rusa* sp.) and the Calamian hog deer (*Axis calamianensis*). Morphometric and morphological traits of antlers, teeth and post-crania were used to distinguish the two cervid species, as reported by Ochoa (2019). The large cervid (*Rusa* sp.) is the more abundant of the two species. Fossil measurements for the Palawan *Rusa* fall at the upper end of the size range of *R. marianna*, whose modern conspecifics have body weights ranging from 40–96kg (see Ochoa & Piper 2017; Ochoa 2019). *Axis calamianensis* is a much smaller deer, with modern individuals weighing 23–40kg. The LGM faunal assemblage is dominated by long bone shaft fragments and axial skeleton specimens of large mammals; although these cannot be definitively ascribed to taxon, most specimens are in the size category of the large cervid (*Rusa* sp.). All other taxa occur in much smaller numbers; nonetheless, these confirm the LGM presence and support the native status of these species on Palawan. This includes rare fossil finds ( $n = 9$ ) for the tiger (*Panthera tigris*).

The tiger's presence may call into question the anthropogenic origin of the assemblage. Only 13 bone fragments (0.16 per cent), however, display carnivore gnawing marks. Based on various butchery and heat-induced modifications, the overall taphonomic evidence suggests that the LGM vertebrate assemblage is primarily human-derived (Table 3). Skeletal element representation in the Layer 2 midden (contexts 111, 112 and 117) indicates that all

Table 3. Summary of vertebrate taphonomic data for Pilanduk Cave, expressed as percentages of total number of bone fragments (TNF) per archaeological context in Trench 3. Weathered = surface weathering; Mn = presence of manganese oxide staining; burnt = charred and calcined bones; gnawed = presence of animal gnawing; Cutmarks = presence of cutmarks; TNF long bone = total number of long bone fragments; helical fracture = long bone fragments with helical fractures.

Context	TNF	% weathered	% Mn	% burnt	% gnawed	% cutmarks	TNF long bone	% helical fracture
111	2174	12.1	11.1	45.8	<0.5	0.7	688	67.4
112	2322	3.4	4.8	39.9	<0.5	0.6	795	59.5
117	2561	0.5	0.9	46.2	<0.5	1.5	1183	17.0
118	506	1.2	0.4	27.9	<0.5	0	221	8.1

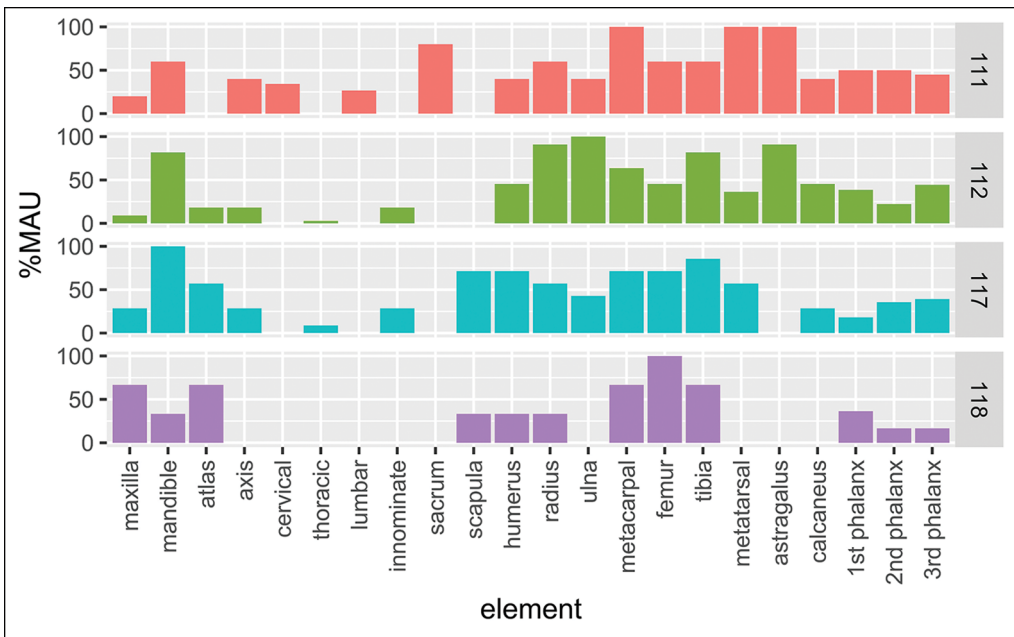


Figure 5. Skeletal element frequencies for deer in Pilanduk Cave, shown as %MAU (minimal animal unit) for Trench 3 archaeological contexts. (© J. Ochoa).

deer body parts are represented in the assemblage (Figure 5). Butchery marks observed on cervid and large mammal remains include helical fractures on long bone fragments, impact scars, cut marks and chop marks (Table 3; Figure S3). Heat alteration on bones is also observed, with a high percentage (40–46 per cent) of bones from Layer 2 displaying charring and calcination (Table 3). The taphonomic evidence suggests that whole deer carcasses were brought in, consumed and discarded by human occupants of the cave, and that the LGM occupation levels of this site were used for deer processing and consumption (Ochoa 2019).

Table 4. Number of identified specimens (NISP) for molluscs from the 2016 Pilanduk excavation, aggregated by habitat ecology. Uppermost layers are mixed, but most specimens are from Layer 2 midden deposits. For additional NISP data, see Tables S2 and S3 in the OSM.

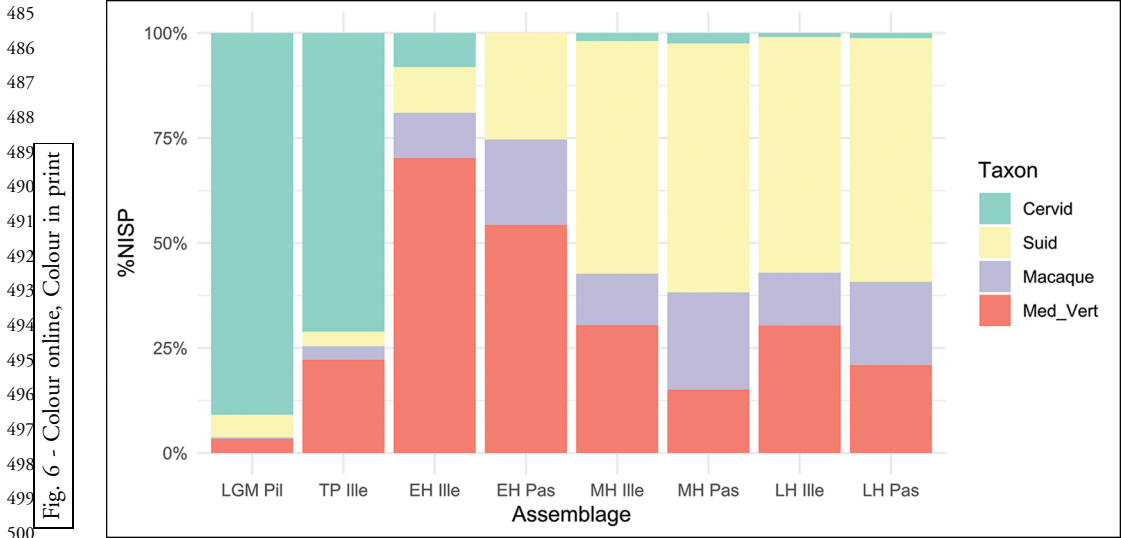
All trenches (1–4)				
Layer	NISP by ecology			
	Terrestrial	Freshwater	Marine	Mangrove
1 and 2 (mixed)	56	410	24	5
2	64	1682	25	2
3	9	50	1	0
<b>Total</b>	129	2142	50	7
<b>%NISP</b>	5.5	92	2.1	0.3

## Mollusc assemblage

A total of 2461 shell artefacts and fragments were analysed from the four trenches of the 2016 excavation, of which 2329 (95 per cent) were identified to taxon (Tables 4 & S2). Although a few shell artefacts from the uppermost layer appear to belong to the jar-burial assemblage, most of the mollusc remains derive from the LGM midden deposits (Layers 2 and 3). A total of 24 taxa were identified (Table S2). Freshwater mollusc taxa predominate, comprising 92 per cent of the entire assemblage (Table 4). Most of the freshwater gastropods have broken tips, possibly indicating consumption. A similar pattern of species composition was reported by Kress (2000) and this reflects the inland riverine environment of the site used by its LGM occupants. Marine and estuarine molluscs are found in small percentages in the LGM midden, indicating the procurement and long-distance transport of these shells into the Palawan interior.

## Discussion

The Pilanduk Cave record shows inland foraging strategies primarily focused on deer hunting and the procurement of freshwater molluscs. The LGM occupation of the cave occurred during a period when Greater Palawan was at its maximum extent and the site was much further inland (Figure 1). To date, Pilanduk provides the only substantive subsistence record known for the LGM for the entire Philippine archipelago. This record can be contextualised within the wider Palawan archaeological context to provide an approximately 40 000-year sequence of changing environments and subsistence patterns, from the MIS-3 record of Tabon Cave (Choa 2018) to the Terminal Pleistocene and Holocene records of Ille and Pasimbahan Caves (Lewis *et al.* 2008; Ochoa *et al.* 2014). Across this sequence, we observe modern human occupation of the following insular environments: tropical rainforests during MIS 3 (based on the Tabon record); open savanna during MIS 2 (Pilanduk record) and the Terminal Pleistocene (Ille record); and tropical rainforests once more during the Holocene (Ille and Pasimbahan records). Environmental reconstructions for Tabon indicate the presence of closed forest at *c.* 39 000 cal BP (Choa 2018), and lithic residue analysis indicates that the cave's



501 *Figure 6. Stacked percentage plot of NISP (number of identified specimens) counts for cervids, suids, macaques and other medium-sized vertebrates (Med\_Vert) in the Palawan record, aggregated by temporal period from Pilanduk (Pil), Ille and Pasimbahan (Pas) Caves. The 'Med\_Vert' category includes terrestrial and arboreal rainforest mammal and reptile taxa. LGM = Last Glacial Maximum; TP = Terminal Pleistocene; EH = Early Holocene; MH = Middle Holocene; LH = Late Holocene. Ille and Pasimbahan faunal data are from Ochoa 2009 and Ochoa et al. 2014 (© J. Ochoa).*

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508 occupants employed foraging strategies that utilised tropical rainforest plants (Xhaufflair *et al.*  
509 2020).

510 In the LGM record of Pilanduk, specialised hunting of deer is observed, and this interior  
511 lowland foraging strategy appears to have persisted until the Terminal Pleistocene *c.* 14 000  
512 years ago, as evidenced by the Ille record (Figure 6). During the Holocene, notable shifts in  
513 subsistence practices are observed in the Ille and Pasimbahan records (Figure 6), wherein  
514 arboreal and other medium-sized forest game species became more abundant, utilisation  
515 of coastal and marine resources increased as a reflection of the changing landscape, wild  
516 pig replaced deer as the main large-game prey and deer became increasingly rare until their  
517 extinction (Lewis *et al.* 2008; Robles *et al.* 2015; Ochoa & Piper 2017). The LGM data  
518 from Pilanduk stands in contrast to this Holocene record, which exhibits a broad-spectrum  
519 subsistence regime adapted to tropical rainforest habitats in response to environmental  
520 changes during the Pleistocene–Holocene transition.

521 The savanna environment reported for Palawan during the LGM (Bird *et al.* 2007, Wurster  
522 *et al.* 2010) was possibly more suitable for the three locally extinct large mammals of the island  
523—the tiger and two deer species. In turn, Pilanduk was occupied by humans focused on hunt-  
524 ing large game prey, particularly the *Rusa* deer. The last fossil occurrence for the Palawan *Rusa* is  
525 from the Early Holocene, while the hog deer is last recorded during the Late Holocene (Ochoa  
526 & Piper 2017). The hog deer now only retains a relict distribution in the Calamianes, a small  
527 cluster of islands immediately north of Palawan (Figure 1). We do not yet have ecological infor-  
528 mation for the Palawan *Rusa*; however, the Calamian hog deer is known to prefer grasslands and

529 open woodlands (Widmann & Lastica 2015). At other sites across Southeast Asia, drier con-  
530 ditions during the LGM led to abandonment and relocation to other locales (O'Connor & Bul-  
531 beck 2014); in the case of Pilanduk, the cave served as an LGM occupation site, abandoned  
532 thereafter, and reused as a jar-burial site much later, in the Late Holocene.

533 What stands out in the records for Palawan and Island Southeast Asia are the multiple  
534 niches that modern humans have occupied—a diverse corpus of adaptations dubbed by Rob-  
535 erts and Stewart (2018) as the ‘generalist specialist’ niche. This hypothesis underscores the  
536 ecological plasticity of Pleistocene *H. sapiens* populations, who colonised a diversity of habi-  
537 tats (‘generalist’) and also developed specialised adaptations to environmental extremes or to  
538 specific ecosystems (‘specialist’). Firstly, where tropical rainforests are present in Island South-  
539 east Asia, *H. sapiens* appears to have utilised successfully a wide range of rainforest resources.  
540 Such broad-spectrum foraging economies are evident at lowland sites where rainforest habi-  
541 tats were present, from the earliest phases of modern human colonisation of the region (*c.* 70  
542 000–50 000 years ago; Barker *et al.* 2017; Westaway *et al.* 2017) to the Holocene, when rain-  
543 forests expanded once more. Secondly, the continuous occupation of the region also neces-  
544 sitated adaptations to other ecological niches in the face of drastic environmental change.  
545 Where grassland/savanna environments expanded during the LGM, we observe shifts in for-  
546 aging behaviour that targeted specific resources, as exemplified by the Pilanduk record.  
547 Thirdly, maritime occupation in Island Southeast Asia also required adaptation to a more  
548 limited terrestrial faunal resource base, and a reliance on coastal, estuarine and marine  
549 resources (Hawkins *et al.* 2017; Shipton *et al.* 2020). Within this context, the Palawan arch-  
550 aeological sequence clearly attests and lends long-term detail to the ecological plasticity attrib-  
551 uted to modern human foragers. It also contributes to our understanding of the versatile and  
552 shifting adaptations of the foragers that have colonised the insular tropics of Southeast Asia.

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## Supplementary information

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